A THEORETICAL TREATMENT OF MICROSCOPIC PHENOMENA IN POROUS ROCKS

by

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Abstract

This thesis makes five significant contributions to the theoretical analysis of the connection between microscopic phenomena and the macroscopic physical properties of porous rocks.

A random sphere packing model permitting contact generation during hydrostatic compression is derived. It is demonstrated that the closure of near-contact gaps with an extremely small mean width significantly alters the elastic properties of granular media from those predicted by previous models in the pressure range used in laboratory measurements.

Generalized forms of the inclusion-based formulations are obtained; major classes of these formulations are defined according to the manner in which interactions between inclusions in the heterogeneous system are simulated. Each class possesses an associated microstructure that determines the topological relationship between the various components.

Inclusion-based formulations are obtained that describe the effects of pore-scale fluid distribution on the dielectric and elastic responses of a partially saturated rock. It is found that the pore fluid configuration within the individual pores and the pore geometries in which saturation conditions are varying are critical factors in determining the dielectric and elastic properties of partially saturated rocks. In addition, it is observed that the effect of saturation condition variations in a particular pore geometry increases as the pore shape becomes more crack-like.
The theoretical formulations describing the effects of pore-scale fluid distribution are used to analyze experimental data for a partially saturated tight gas sandstone and glass bead packing. Simple models that incorporate only the basic geometrical elements of the resulting pore-scale fluid distributions accurately predict the experimental data. It is also found that the same geometrical model can be used simultaneously to estimate the dielectric and elastic responses of a partially saturated porous medium.

The effect of surface phenomena (e.g. electrical double layers and surface conduction) at the solid-fluid interface are incorporated into inclusion-based formulations of the dielectric response by employing the limiting case of a confocally-layered ellipsoid. It is found that the effect of surface phenomena varies as a function of both inclusion size and shape.
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Chapter One

Prologue

"The book of nature is written in the language of mathematics."

Galileo Galilei (1564-1642)

Geophysics, as a discipline, is frequently concerned with the problem of determining the nature of porous rocks that cannot be directly observed. To achieve this goal, geophysical surveys are conducted to measure the spatial and/or temporal distribution of particular physical properties. However, these physical property measurements provide only an indirect means of determining and analyzing microscopic phenomena in porous rock. When one is concerned with these types of phenomena, it is necessary to establish an appropriate framework in which geophysical measurements can be interpreted.

One approach is to perform laboratory measurements on porous rock samples that are believed to be similar to those that exist in situ. However, these types of measurements can prove to be of limited utility. Minor differences in terms of rock type and/or the conditions to which the rock is subjected can result in significant variations in physical behavior. Therefore, it is necessary to complement these laboratory results by developing mathematical formulations that describe the physical response of porous rocks in terms of its microscopic phenomena. In this study, theoretical formulations are developed which describe the role of certain microscopic phenomena in determining the observed physical
properties of porous rocks. Given the relative importance of electromagnetic and seismic techniques in geophysics, this study will be confined to the consideration of dielectric and elastic response of porous rocks.

In unconsolidated sands, an important type of microscopic phenomenon is the interaction between neighboring sand grains. Experimental data for granular materials consistently violate the relationship between elastic wave velocities and confining pressure predicted by previous sphere packing models. Murphy (1982a) proposed that this discrepancy between these earlier models and the experimental results is due to the creation of grain contacts as confining pressure increases. A random sphere packing model that permits the generation of grain contacts is derived in Chapter Two. However, this type of formulation is specifically structured to incorporate interactions between neighboring particles. Hence, it is not well suited for the study of phenomena that are related to geometrical configuration of the constituents within a porous rock.

In the range of frequencies of interest in the geophysical measurement of dielectric behavior (i.e. 100 kHz to 1 GHz), a dominant mechanism affecting the dielectric response of porous rocks is the charge polarization at interfaces on the microscopic scale (Poley et al., 1978). Further, it has been observed that the geometry of the pore spaces within a rock significantly affects its elastic wave velocities (Nur and Simmons, 1969; Toksoz et al., 1976). Therefore, it is necessary to use a formulation that incorporates sufficient information about the pore space geometries within the system to describe elastic and electromagnetic wave propagation in porous rocks. This is accomplished by employing an inclusion-based formulation to describe porous rocks. In this study, the porous rock system is viewed as a rock matrix into which inclusions representing the individual pore
spaces are embedded. In previous work, a diverse collection of mathematical expressions have been obtained when the inclusion-based formulation has been employed to describe a given physical property of composite materials. To clarify the relationship between these formulations, the major classes of the inclusion-based formulations are derived in Chapter Three in such a way as to avoid special cases that result from the use of specific inclusion types. The discussion is in terms of dielectric properties; however, the basic methods and principles used in the derivation are applicable to inclusion-based formulations for other physical properties, such as elastic wave propagation.

A phenomenon that is characterized by the geometrical configuration of the constituents within a porous rock is the pore-scale fluid distribution within the partially-saturated system. In Chapter Four, mathematical expressions are developed for the dielectric and elastic response of partially saturated porous rocks for various geometrical configurations of the pore fluids within the individual pore space. To determine the critical factors that control the relationship between pore-scale fluid distribution and macroscopic physical properties, a number of numerical simulations are performed utilizing these formulations.

The basic results obtained from this theoretical study are applied in Chapter Five to the analysis of experimental data collected on partially saturated porous rocks. The measured apparent dielectric constant and elastic wave velocities for a tight gas sandstone undergoing an imbibition-drainage cycle (Knight and Nur, 1987a; Knight and Nolen-Hoeksema, 1990) are modelled by means of simple geometrical scenarios for the resulting pore-scale fluid distribution. In addition, elastic wave velocity data for a
partially saturated glass bead packing obtained by Domenico (1977) are examined. In the derivation of the inclusion-based formulations, the existence of surface phenomena at the fluid-solid interface has been explicitly neglected, although such phenomena can have a significant effect on the electrical properties of porous rock. Previously, these phenomena were implicitly incorporated into mathematical formulations by assigning their effects to one of the constituents and adjusting its properties accordingly (Knight and Endres, 1990). However, this approach does not preserve the geometrical structure inherent in the system. In Chapter Six, an inclusion-based formulation is presented that explicitly incorporates the effects of surface phenomena in the dielectric response of a heterogeneous medium. This is achieved by taking a limiting case of a confocally-layered ellipsoid as its outer shell becomes infinitesimally thin.
Chapter Two

A Sphere Packing Model Incorporating Contact Generation

Introduction

The modelling of granular materials by means of sphere packings has been the subject of numerous papers. Many authors (Gassmann (1951a), Duffy and Mindlin (1957), Deresiewicz (1958), White (1965), and Walton (1975)) employed regular packings (e.g. hexagonal close packing, cubic packing). Brandt (1955) presented the first formulation for the elastic behavior of random packings; however, only the overall volumetric changes and the effective bulk modulus of the medium were considered. More recently, Digby (1981) and Walton (1987) have presented more comprehensive treatments for the mechanical behavior of random sphere packings. In all of these studies, it was assumed that grain contacts were established in the initial, undeformed state and that no new grain contacts were generated during the compression of the medium. Further, some specific form of contact microphysics was always assumed.

Experimental evidence has shown that these mathematical models do not adequately describe the observed elastic behavior of sphere packings. In all of the above models, the hydrostatic strain and the effective elastic moduli are shown to be dependent on the $2/3$ and $1/3$ power, respectively, of the hydrostatic confining pressure. This implies that the elastic wave velocities are related to confining pressure by essentially a $1/6$ power law. Duffy and Mindlin (1957) presented experimental data for the elastic wave velocities of a
face-centered cubic packing of spheres. These data showed that the predicted elastic wave velocity was significantly higher than the measured velocity, and did not have a $1/6$ power dependence on confining pressure. Further, as the allowable variation in sphere radii increased, the deviation between predicted and experimental velocity also increased. Domenico (1977) measured the elastic wave velocities of the random packing of air saturated glass beads under hydrostatic confining pressures between 2.7 and 34.4 MPa. A least squares fit to the data implied that both compressional and shear wave velocities have approximately a $1/4$ power relationship with the confining pressure. Murphy (1982a) measured the compressional and shear wave velocities of a random packing of glass beads under vacuum and hydrostatic confining pressures between 0.1 and 35 MPa. A $1/4$ power relationship was observed at low confining pressures, and a $1/6$ power relationship was exhibited at high confining pressures.

Murphy (1982a) stated that the $1/4$ power relationship can be explained by the generation of contacts due to the closing of near-contact gaps and used a simple two-layer model to illustrate this point. However, a general formulation for a random packing was not derived. In this chapter, expressions are derived for the overall stress-strain relationship and the effective moduli of a granular medium under hydrostatic compression, which permit the generation of grain contacts due to the closure of near-contact gaps. The mathematical formulation of Walton (1987) provides a basis which can be expanded to include the effect of contact generation caused by the closing of near-contact gaps as the hydrostatic pressure increases. Further, Walton (1987) considers two cases of contact microphysics (i.e. perfectly smooth and infinitely rough spheres). The formulation presented here extends these results by allowing the use of general contact...
microphysics.

Undefomed Configuration

Consider a random packing of identical solid spheres of radius \( R \). Each sphere is elastically isotropic and homogeneous with Lamé constants \( \mu \) and \( \lambda \). Let \( V \) be a volume within the packing containing \( N \) spheres and having dimensions which are very large compared to \( R \). Within \( V \), the random packing is assumed to be homogeneous.

Let \( x^{(n)} \) be the position of the center of the \( n \)th sphere relative to some given Cartesian coordinate system. The \( n \)th sphere is surrounded by neighboring spheres, some of which are in point contact with the \( n \)th sphere, while others are in near-contact, as in Figure (2.1). In Figure (2.2), the system of two spheres initially in point contact (in this case, the \( m \)th and \( n \)th spheres) is shown. Let the midpoint between sphere centers be denoted by \( O \), with a position vector \( x^{(o)} \) given by

\[
{x^{(o)}} = \frac{1}{2} \left( x^{(n)} + x^{(m)} \right). \tag{2.1}
\]

For this system, the inward unit vector along the line of the sphere centers, \( i^{(nm)} \), is defined by

\[
i^{(nm)} = (2R)^{-1} \left( x^{(n)} - x^{(m)} \right). \tag{2.2}
\]

The system of two spheres in near-contact is illustrated in Figure (2.3). In this case, let \( O' \)
be the midpoint between the sphere centers, with a position vector \( \mathbf{x}^{(o')} \) given by

\[
\mathbf{x}^{(o')} = \frac{1}{2} \left( \mathbf{x}^{(n)} + \mathbf{x}^{(m')} \right).
\] (2.3)

Similarly, the inward unit vector along the line of the sphere centers, \( \mathbf{i}^{(nm')} \), for this system is defined by

\[
\mathbf{i}^{(nm')} = \left[ 2 \left( \mathbf{R} + \mathbf{g}^{(nm')} \right) \right]^{-1} \left( \mathbf{x}^{(n)} - \mathbf{x}^{(m')} \right),
\] (2.4)

where \( 2\mathbf{g}^{(nm')} \) is the distance (or gap) between the surface of \( n \)th and \( m' \)th spheres along the line of the sphere centers.

**Overall Stress-Strain Relationship**

In order to derive the overall stress-strain relationship, consider an initial deformed state, obtained by subjecting the boundary of volume \( V \) to a displacement \( \mathbf{u} \) consistent with a uniform, compressive hydrostatic strain \( \varepsilon \mathbf{I} \) (where \( \mathbf{I} \) is the identity tensor). Hence, for a position vector \( \mathbf{x} \) in \( V \), the displacement at this point if \( V \) were a continuum is given by

\[
\mathbf{u} (\mathbf{x}) = \varepsilon \mathbf{I} \cdot \mathbf{x} = \varepsilon \mathbf{x}.
\] (2.5)

It will be assumed that the sphere centers undergo a deformation which is consistent with this uniform strain and that all deformations are reversible. Then,
\[ u^{(n)} = \varepsilon x^{(n)} \]  

(2.6)

defines the displacement of the \( n \)th sphere center. Similar expressions give the 
displacements for the centers of the \( m \)th and \( m' \)th sphere.

Consider the situation where two spheres are in point contact in the undeformed state. 
If \( \Delta^{(no)} \) is the displacement of midpoint O with respect to the \( n \)th sphere center, then

\[ \Delta^{(no)} = \frac{1}{2} (u^{(m)} - u^{(n)}) = -R \varepsilon i^{(nm)}. \]  

(2.7)

As the hydrostatic strain increases, the contact area between spheres \( n \) and \( m \) becomes 
finite; and, a force, \( F^{(nm)} \), is exerted on the \( n \)th sphere due to its contact with the \( m \)th. 
Since the spheres are elastically isotropic and homogeneous, it can be assumed that \( F^{(nm)} \) 
is parallel to \( i^{(nm)} \). Hence,

\[ F^{(nm)} = F_N (\Phi^{(nm)}) i^{(nm)}, \]  

(2.8)

where \( F_N \) is the magnitude of the normal component of \( F^{(nm)} \) and \( \Phi^{(nm)} \) defines the state 
of the contact between the \( n \)th and \( m \)th spheres. Since all original contacts undergo 
identical deformations for a given hydrostatic strain \( \varepsilon \), \( \Phi^{(nm)} = \Phi \) for all these contacts.

For the spheres in near-contact, it is first necessary to establish a point contact before 
a force is exerted on the \( n \)th sphere. Since the relative displacement of one sphere center 
with respect to another is along the line of their centers, it can be shown that the point
contact is obtained when

\[ \| \Delta^{(no')} \| = g^{(nm')} , \]  

(2.9)

where

\[ \Delta^{(no')} = - (R + g^{(nm')}) \varepsilon \ i^{(nm')} , \]  

(2.10)

\( \Delta^{(no')} \) being the displacement of the midpoint \( O' \) with respect to the \( n \)th sphere center.

Once this occurs, a finite contact area is established as the hydrostatic strain continues to increase; and, a force, \( F^{(nm')} \), is exerted on the \( n \)th sphere due to its contact with the \( m' \)th sphere. Given the above assumptions, the force \( F^{(nm')} \) is given by

\[ F^{(nm')} = F_N \left( \Phi^{(nm')} \right) H \left( \| \Delta^{(no')} \| - g^{(nm')} \right) i^{(nm')} \]

\[ = F_N \left( \Phi^{(nm')} \right) H \left( - (R + g^{(nm')}) \varepsilon - g^{(nm')} \right) i^{(nm')} , \]  

(2.11)

where \( \Phi^{(nm')} \) defines the state of the contact between spheres \( n \) and \( m' \), and \( H \) is the Heaviside step function.

Now that the forces on the \( n \)th sphere due to its contacts with its neighboring spheres are known, it is possible to relate the average stress within \( V \) to the average strain. The average stress, \( \langle \sigma_{ij} \rangle \), within \( V \) is given by
\[ \langle \sigma_{ij} \rangle = \frac{1}{V} \int_{\text{spheres}} \sigma_{ij} \, dV = \frac{1}{V} \sum_n \int_{V_n} \sigma_{ij}^{(n)} \, dV , \]  

(2.12)

where \( V_n \) is the volume of the \( n \)th sphere, \( \sigma_{ij}^{(n)} \) are the components of the Cauchy stress within the \( n \)th sphere, and the summation is over all spheres in \( V \). For the \( n \)th sphere, the stress within \( V_n \) can be related to the tractions on \( S_n \), the surface of the \( n \)th sphere, in the following manner:

\[ \int_{V_n} \sigma_{ij}^{(n)} \, dV = \frac{1}{2} \int_{S_n} \left( x_i' t_{ij}^{(n)} + x_j' t_{ij}^{(n)} \right) \, ds , \]  

(2.13)

where \( t^{(n)} \) is the traction on \( S_n \) and \( x' = x - x^{(n)} \) is the position of a point on \( S_n \) with respect to the center of the \( n \)th sphere. In general, it can be assumed that the contact areas will be small. This allows two simplifying assumptions. First, \( x' \) may be approximated by

\[ x' = \frac{1}{2} \left( x^{(m)} - x^{(n)} \right) = - R \, i^{(nm)} \]  

(2.14)

for the original contacts and by

\[ x' = \frac{1}{2} \frac{R}{R + g^{(nm')}} \left( x^{(m')} - x^{(n)} \right) = - R \, i^{(nm')} \]  

(2.15)

for the generated contacts. Secondly, the integration of the \( i \)th component of the traction
over the contact surface is $F^{(nm)}_i$ for an original contact and $F^{(nm')}_i$ for the generated contacts.

Using these approximations, Equation (2.13) becomes

$$
\int_{V_n} \sigma_{ij}^{(n)} \, dV = \frac{-R}{2} \times 
$$

$$
\left\{ \sum_m \left( i_{i}^{(nm)} F_{j}^{(nm)} + i_{j}^{(nm)} F_{i}^{(nm)} \right) + \sum_m \left( i_{i}^{(nm')} F_{j}^{(nm')} + i_{j}^{(nm')} F_{i}^{(nm')} \right) \right\}. \quad (2.16)
$$

In this expression, the first term of the right hand side is the contribution of the original contacts and the second term represents the effect of created contacts. By using Equation (2.16) in Equation (2.12) and changing the summation from over the indices to over individual contacts, the following is obtained

$$
\langle \sigma_{ij} \rangle = \frac{-R}{V} \left\{ \sum_{\text{orig. cont.}} \left( i_{i}^{(nm)} F_{j}^{(nm)} + i_{j}^{(nm)} F_{i}^{(nm)} \right) + \right. 
$$

$$
\left. \sum_{\text{creat. cont.}} \left( i_{i}^{(nm')} F_{j}^{(nm')} + i_{j}^{(nm')} F_{i}^{(nm')} \right) \right\}. \quad (2.17)
$$

The fact that the medium is undergoing hydrostatic compression is now introduced.

This means

$$
\langle \sigma_{ij} \rangle = -p \delta_{ij}, \quad (2.18)
$$
where \( p \) is the pressure applied to the surface of \( V \) and \( \delta_{ij} \) is the Kronecker delta. By employing Equations (2.8), (2.11), and (2.18), Equation (2.17) becomes

\[
p \delta_{ij} = \frac{2R}{V} \left\{ \sum_{\text{orig. cont.}} F_N(\Phi) \langle i_i^{(nm)} i_j^{(nm)} \rangle + \right. \\
\left. \sum_{\text{creat. cont.}} F_N(\Phi^{(nm')}) \left[ - (R + g^{(nm')}) \epsilon - g^{(nm')} \right] \langle i_i^{(nm')}, i_j^{(nm')} \rangle \right\}. \tag{2.19}
\]

Since \( V \) contains a large number, \( N \), of spheres, Equation (2.19) can be expressed in terms of average values for the type of contact involved as follows:

\[
p \delta_{ij} = \frac{RN}{V} \left\{ \eta_0 F_N(\Phi) \langle i_i^{(nm)} i_j^{(nm)} \rangle + \right. \\
\left. \eta_c \langle F_N(\Phi^{(nm')}) \left[ - (R + g^{(nm')}) \epsilon - g^{(nm')} \right] \rangle \langle i_i^{(nm')}, i_j^{(nm')} \rangle \right\}, \tag{2.20}
\]

where \( \eta_0 \) and \( \eta_c \) are the average number of original contacts per sphere and the average number of near-contacts per sphere, respectively, and \( \langle \rangle \) represent average values taken over the contacts. Since it has been assumed that the packing is statistically homogeneous and isotropic on the scale of the dimensions of \( V \), then original and created contact points have a uniform probability distribution over a sphere surface. In this

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situation, Batchelor and O'Brien (1977) show that

\[
\langle i_i^{(nm)} i_j^{(nm)} \rangle = \langle i_i^{(nm')} i_j^{(nm')} \rangle = \frac{1}{3} \delta_{ij} .
\] (2.21)

Further, using

\[
V = \frac{4\pi R^3 N}{3(1 - \phi)},
\] (2.22)

where \(\phi\) is the porosity of the undeformed aggregate, Equation (2.20) becomes

\[
p = \left(1 - \phi\right) \left[ \eta_o F_N(\Phi) + \eta_c \left\{ F_N(\Phi^{(nm')}) H(\cdot (R + g^{(nm')}) \epsilon - g^{(nm')}) \right\} \right].
\] (2.23)

The average value of the normal force at the created contacts can be determined if the probability density function of \(g^{(nm')}\), \(f(g^{(nm')})\), is known. Since \(g^{(nm')} \geq 0\) and the Heaviside step function is zero if

\[
g^{(nm')} \geq -R \epsilon (1 + \epsilon)^{-1},
\] (2.24)

then, Equation (2.23) can be written as

\[
p = \left(1 - \phi\right) \left[ \eta_o F_N(\Phi) + \eta_c \int_0^\infty F_N(\Phi^{(nm')}) f(g^{(nm')}) dg^{(nm')} \right].
\] (2.25)
In order to obtain the overall stress-strain relationship, the form of the contact microphysics must be defined. A specific example is considered in the section on numerical results below.

**Effective Moduli**

In order to determine the effective moduli of the granular medium, it is necessary to consider the effect of the application of an additional incremental deformation to the sphere packing after the initial hydrostatic deformation. Therefore, assume that the boundary of $V$ is subjected to a further deformation $\delta u$ which is consistent with a uniform strain tensor $\delta E$. Then, the displacement of a point $x$ in $V$, if it were in a continuum, would be given by

$$\delta u(x) = \delta E \cdot x, \quad \text{or} \quad \delta u_i(x) = \delta \epsilon_{ij} x_j. \quad (2.26)$$

It will be assumed that this incremental deformation is small enough that effects of contact generation and destruction can be neglected, that the sphere centers undergo an incremental deformation which is consistent with a uniform incremental strain, and that all incremental deformations are reversible. Hence,

$$\delta u^{(n)} = \delta E \cdot x^{(n)}, \quad \text{or} \quad \delta u_i^{(n)} = \delta \epsilon_{ij} x_j^{(n)}. \quad (2.27)$$

Given this as a starting point, then the incremental relative displacements of the midpoints $O$ and $O'$ with respect to the $n$th sphere center are given by
\[ \delta \Delta^{(no)} = -R \delta E \cdot i^{(nm)} \]  

(2.28)

and

\[ \delta \Delta^{(no') } = -(R + g^{(nm') }) \delta E \cdot i^{(nm')}. \]  

(2.29)

Since the spheres are composed of homogeneous, isotropic material, it can be assumed

that the incremental force at an original contact is given by

\[ \delta F^{(nm)} = \delta F_N(\Phi) i^{(nm)} + \delta F_T(\Phi) j^{(nm)}, \]  

(2.30)

where

\[ j^{(nm)} = \frac{-R \left[ \delta E \cdot i^{(nm)} - \delta E_{kl} i_k^{(nm)} i_l^{(nm)} i^{(nm)} \right]}{|| R \left[ \delta E \cdot i^{(nm)} - \delta E_{kl} i_k^{(nm)} i_l^{(nm)} i^{(nm)} \right] ||}. \]  

(2.31)

Similarly, the incremental force at a created contact is

\[ \delta F^{(nm')} = \left[ \delta F_N(\Phi^{(nm')}) i^{(nm')} + \delta F_T(\Phi^{(nm')}) j^{(nm')} \right] \times \]

\[ H \left( -(R + g^{(nm')}) \varepsilon - g^{(nm')} \right), \]  

(2.32)

where
\[
\mathbf{j}^{(\text{nm}')} = \frac{-\left(\mathbf{R} + g^{(\text{nm}')}\right) \left[ \mathbf{\delta E} \cdot \mathbf{i}^{(\text{nm}')} - \left(\mathbf{\delta e}_{kl} i_k^{(\text{nm}')} i_l^{(\text{nm}')}\right) \mathbf{i}^{(\text{nm}')} \right]}{|| - \left(\mathbf{R} + g^{(\text{nm}')}\right) \left[ \mathbf{\delta E} \cdot \mathbf{i}^{(\text{nm}')} - \left(\mathbf{\delta e}_{kl} i_k^{(\text{nm}')} i_l^{(\text{nm}')}\right) \mathbf{i}^{(\text{nm}')} \right] ||},
\]

and \(\delta F_N\) and \(\delta F_T\) are the magnitude of the normal and tangential component of the incremental contact force \(\delta F\), respectively.

The concept of the contact stiffness is now introduced. Let \(D_N\) and \(D_T\) be the normal and tangential contact stiffness, respectively. Then,

\[
D_N = \frac{dF_N}{dw},
\]

and

\[
D_T = \frac{dF_T}{dr},
\]

where \(w\) is the magnitude of the normal component of the relative displacement of \(O\) or \(O'\) (i.e. parallel to \(i^{(\text{nm})}\) or \(i^{(\text{nm}')}\)) and \(r\) is the magnitude of the tangential component of the relative displacement of \(O\) or \(O'\) (i.e. perpendicular to \(i^{(\text{nm})}\) or \(i^{(\text{nm}')}\)). Since the incremental displacement is small, then

\[
\delta F_N = D_N \delta w,
\]

and

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\[ \delta F_T = D_T \delta r. \] (2.37)

Hence, the contact forces can be rewritten as

\[ \delta F^{(nm)} = - R \left\{ \left( D_N(\Phi) - D_T(\Phi) \right) \left( \delta \varepsilon_{kl} i_k^{(nm)} i_{l_i}^{(nm)} \right) i_j^{(nm)} + D_T(\Phi) \delta E \cdot i^{(nm)} \right\} \] (2.38)

for the original contacts, and

\[ \delta F^{(nm')} = - (R + g^{(nm')}) \left\{ \left( D_N(\Phi^{(nm')}) - D_T(\Phi^{(nm')}) \right) \left( \delta \varepsilon_{kl} i_k^{(nm')} i_{l_i}^{(nm')} \right) i_j^{(nm')} + \right. \]

\[ D_T(\Phi^{(nm')}) \delta E \cdot i^{(nm')} \right\} H \left\{ -(R + g^{(nm')} e - g^{(nm')}) \right\} \] (2.39)

for the created contacts.

Using the same procedure as was used to determine the relationship between pressure and hydrostatic strain, the following relationship is found between the average incremental stress, \( \langle \delta \sigma_{ij} \rangle \), and the incremental average strain \( \langle \delta \varepsilon_{kl} \rangle \):

\[ \langle \delta \sigma_{ij} \rangle = \frac{3(1 - \phi)}{8\pi R^2} \left\{ \eta_0 R \left[ 2 \left( D_N(\Phi) - D_T(\Phi) \right) \langle \delta \varepsilon_{kl} \rangle \left( i_k^{(nm)} i_{l_j}^{(nm)} i_j^{(nm)} i_{l_i}^{(nm)} \right) + \right. \right. \]

\[ \left. \left. D_T(\Phi) \left( \langle \delta \varepsilon_{jk} \rangle \left( i_j^{(nm)} i_{l_k}^{(nm)} \right) + \langle \delta \varepsilon_{ik} \rangle \left( i_i^{(nm)} i_{l_j}^{(nm)} \right) \right) \right] + \eta_e \times \right. \]
\[
\begin{align*}
2 \left( \langle D_N \Phi^{(nm')} - D_T \Phi^{(nm')} \rangle \right) (R + g^{(nm')}) H \left( - (R + g^{(nm')}) \epsilon - g^{(nm')} \right) \langle \delta e_{kl} \rangle \times \\
\langle i^{(nm)}_i j^{(nm)}_j k^{(nm)}_k l^{(nm)}_l \rangle + \langle D_T \Phi^{(nm')} \rangle (R + g^{(nm')}) H \left( - (R + g^{(nm')}) \epsilon - g^{(nm')} \right) \times \\
\left( \langle \delta e_{jk} \rangle \langle i^{(nm)}_k j^{(nm)}_j \rangle + \langle \delta e_{ik} \rangle \langle i^{(nm)}_l j^{(nm)}_j \rangle \right) \right). 
\end{align*}
\]

(2.40)

Since the initial deformation was hydrostatic, the homogeneous and isotropic nature of the random packing is retained during the incremental deformation. Hence, Equation (2.21) and another result from Batchelor and O’Brien (1977),

\[
\langle i^{(nm)}_i j^{(nm)}_j k^{(nm)}_k l^{(nm)}_l \rangle = \langle i^{(nm')}_i j^{(nm')}_j k^{(nm')}_k l^{(nm')}_l \rangle \\
= \frac{1}{15} \left( \delta_{ij} \delta_{kl} + \delta_{ik} \delta_{jl} + \delta_{il} \delta_{jk} \right)
\]

(2.41)

may be used.

By inserting Equations (2.21) and (2.41) into (2.40), the resulting expression can be compared with

\[
\langle \delta e_{ij} \rangle = C_{ijkl}^* \langle \delta e_{kl} \rangle,
\]

(2.42)

where the effective moduli, \( C_{ijkl}^* \), can be expressed in terms of the effective Lamé
constants, $\mu^*$ and $\lambda^*$, as follows

$$C^*_{ijkl} = \lambda^* \delta_{ij} \delta_{kl} + \mu^* (\delta_{ik} \delta_{jl} + \delta_{il} \delta_{jk}).$$  \hspace{1cm} (2.43)

Hence, it can be found that

$$\lambda^* = \left(1 - \phi \right) \left\{ \frac{\eta_o}{R^2} \left( D_N(\Phi) - D_T(\Phi) \right) + \frac{\eta_c}{R^2} \times ight.$$

$$\left( \left[ D_N\left( \Phi^{(nm)} \right) - D_T\left( \Phi^{(nm)} \right) \right] (R + g^{(nm)}) H \left[ - (R + g^{(nm)}) \varepsilon - g^{(nm)} \right) \right) \right\}. \hspace{1cm} (2.44)$$

and

$$\mu^* = \left(1 - \phi \right) \left\{ \frac{\eta_o}{R^2} \left[ D_N(\Phi) + \frac{3}{2} D_T(\Phi) \right] + \frac{\eta_c}{R^2} \times ight.$$

$$\left( \left[ D_N\left( \Phi^{(nm)} \right) + \frac{3}{2} D_T\left( \Phi^{(nm)} \right) \right] (R + g^{(nm)}) H \left[ - (R + g^{(nm)}) \varepsilon - g^{(nm)} \right) \right) \right\}. \hspace{1cm} (2.45)$$

Letting $f(g^{(nm)})$ again be the probability density function of $g^{(nm)}$, then, as was done above, $\lambda^*$ and $\mu^*$ can be written as
\[
\lambda^* = \frac{(1 - \phi)}{20\pi} \left\{ \frac{\eta_0}{R} (D_N(\Phi) - D_T(\Phi)) + \frac{\eta_c}{R^2} \times \right. \\
\left. \int_0^{R \epsilon(1 + \epsilon)^{-1}} \left[ (D_N(\Phi^{(nm')}) - D_T(\Phi^{(nm')}) \right] (R + g^{(nm')}) f(g^{(nm')}) \, dg^{(nm')} \right\}, \tag{2.46}
\]

and

\[
\mu^* = \frac{(1 - \phi)}{20\pi} \left\{ \frac{\eta_0}{R} (D_N(\Phi) + \frac{3}{2} D_T(\Phi)) + \frac{\eta_c}{R^2} \times \right. \\
\left. \int_0^{R \epsilon(1 + \epsilon)^{-1}} \left[ (D_N(\Phi^{(nm')}) + \frac{3}{2} D_T(\Phi^{(nm')}) \right] (R + g^{(nm')}) f(g^{(nm')}) \, dg^{(nm')} \right\}. \tag{2.47}
\]

**Numerical Results**

From the above analysis, it can be seen that the general elastic behavior of a contact generating granular medium is described by Equations (2.25), (2.46), and (2.47). In order to obtain numerical results, it is necessary to define the nature of the contact microphysics and \(f(g^{(nm')})\), the probability density function of \(g^{(nm')}\), for the sphere packing.

The contact microphysics gives the form of the interaction between spheres in contact. Let us assume that spheres are infinitely rough (i.e. no slippage along the contact surfaces). Then, using the results of Walton (1978), the following expressions can be derived for the contact forces and contact stiffness:
\[ F_N(\Phi) = \frac{4R^2 (1 - \varepsilon)^{3/2}}{3\pi B} , \quad (2.48) \]

\[ F_N(\Phi^{(nm')}) = \frac{4R^{1/2}}{3\pi B} (1 - R \varepsilon - g^{(nm')}(1 + \varepsilon))^{3/2} , \quad (2.49) \]

\[ D_N(\Phi) = \frac{2R (1 - \varepsilon)^{1/2}}{\pi B} , \quad (2.50) \]

\[ D_T(\Phi) = \frac{4R (1 - \varepsilon)^{1/2}}{\pi (2B + C)} , \quad (2.51) \]

\[ D_N(\Phi^{(nm')}) = \frac{2[ R (1 - g^{(nm')}(1 + \varepsilon))^{1/2}]}{\pi B} , \quad (2.52) \]

\[ D_T(\Phi^{(nm')}) = \frac{4[ R (1 - g^{(nm')}(1 + \varepsilon))^{1/2}]}{\pi (2B + C)} , \quad (2.53) \]

where

\[ B = \frac{1}{4\pi} \left| \frac{1}{\mu} + \frac{1}{\mu + \lambda} \right| \quad (2.54) \]

and

\[ C = \frac{1}{4\pi} \left| \frac{1}{\mu} - \frac{1}{\lambda + \mu} \right| . \quad (2.55) \]

The nature of the probability density function, \( f(g^{(nm')}) \), has not been established. However, given the nature of the problem, it is plausible that \( g^{(nm')} \) has a Maxwell
distribution. Hence, \( f(g^{(nm)}) \) is given by (Papoulis (1965))

\[
f(g^{(nm)}) = \frac{2^{1/2}}{\alpha^3 \pi^{1/2}} (g^{(nm)})^2 \exp \left[ -(g^{(nm)})^2 (2 \alpha^2)^{-1} \right] H(g^{(nm)}) .
\] (2.56)

The maximum of \( f(g^{(nm)}) \) occurs at \( \alpha \cdot 2^{1/2} \) and has a value of \( 2^{3/2} (\alpha \varepsilon \pi^{1/2})^{-1} \). By using Equations (2.48) through (2.53), and (2.56), the following expressions can be derived for the confining pressure \( p \) and the effective moduli \( \lambda^* \) and \( \mu^* \):

\[
p = \frac{(1 - \phi)(-\varepsilon)^{3/2}}{3\pi^2 B} \left( \eta_o + \eta_e \frac{2^{1/2}}{\alpha^3 \pi^{1/2}} \left( \frac{-R \varepsilon}{1 + \varepsilon} \right)^3 \int_0^1 (1 - x)^{3/2} x^2 \exp(-\beta x^2) \, dx \right),
\] (2.57)

\[
\lambda^* = \frac{(1 - \phi) C (-\varepsilon)^{1/2}}{10\pi^2 B (2B + C)} \left\{ \eta_o + \eta_e \frac{2^{1/2}}{\alpha^3 \pi^{1/2}} \left( \frac{-R \varepsilon}{1 + \varepsilon} \right)^3 \times \right.
\]

\[
\left\{ \int_0^1 (1 - x)^{1/2} x^2 \exp(-\beta x^2) \, dx + \left( \frac{-\varepsilon}{1 + \varepsilon} \right) \int_0^1 (1 - x)^{1/2} x^3 \exp(-\beta x^2) \, dx \right\},
\] (2.58)

\[
\mu^* = \frac{(1 - \phi)(5B + C)(-\varepsilon)^{1/2}}{10\pi^2 B (2B + C)} \left\{ \eta_o + \eta_e \frac{2^{1/2}}{\alpha^3 \pi^{1/2}} \left( \frac{-R \varepsilon}{1 + \varepsilon} \right)^3 \times \right.
\]

\[
\left\{ \int_0^1 (1 - x)^{1/2} x^2 \exp(-\beta x^2) \, dx + \left( \frac{-\varepsilon}{1 + \varepsilon} \right) \int_0^1 (1 - x)^{1/2} x^3 \exp(-\beta x^2) \, dx \right\},
\] (2.59)

where
\[ \beta = (2\alpha^2)^{-1} \left( \frac{-R \varepsilon}{1 + \varepsilon} \right)^2. \]  

(2.60)

The integrals in Equations (2.57) - (2.59) can be shown to reduce to generalized hypergeometric functions (Gradshteyn and Ryzhik (1965)). However, to evaluate these expressions, it is much more practical to perform a numerical integration.

Table (2.1) gives the values of parameters used in evaluating Equations (2.57) - (2.59). These values are typical of those used for dense packings of glass beads in Domenico (1977) and Murphy (1982a). Values for \( \eta_0 \) and \( \eta_c \) were obtained from Bernal and Mason (1960). Figures (2.4), (2.5), and (2.6) illustrate \( \log (-\varepsilon) \), \( \log (\lambda^*) \), and \( \log (\mu^*) \) as functions of \( \log (p) \), respectively. In each case, \( \alpha = 0.001R \) and \( 0.0001R \) were used. In addition, the results for the model of Walton (1987) are also shown for two cases: the number of contacts \( \eta = \eta_0 + \eta_c \) and \( \eta = \eta_0 \).

In each case, the model of Walton (1987) gives linear relationships in terms of \( \log (p) \), corresponding to the 2/3 power law for \( (-\varepsilon) \) and the 1/3 power law for the effective moduli \( \lambda^* \) and \( \mu^* \). From Figures (2.4) - (2.6), it can be seen that three distinct regions exist in the response of the contact generating model. At low pressures, the contact generating model asymptotically approaches Walton's model for \( \eta = \eta_0 \) as \( p \) decreases. At high pressures, the contact generating model asymptotically approaches the completely contacted model, \( \eta = \eta_0 + \eta_c \), as \( p \) increases. Between the two, a transition zone exists. The onset of the transition zone as pressure increases is related to the mean value \( g^{(nm)} \) (i.e. once a significant portion of the possible contacts are established, the transition begins). This behavior of the transition is independent of the choice of the
probability density function, as it is a general response of the sphere packing due to contact generation.

If an attempt were made to describe the response of the model in the transition zone by means of a power law relationship, Figures (2.5) and (2.6) show that the effective exponent would be greater than 1/3 for the relationship between the effective moduli and the confining pressure. Hence, the effective power law between confining pressure and elastic wave velocities would have an exponent greater than 1/6. This is consistent with the experimental results of Domenico (1977) and Murphy (1982a).

Conclusions

In this chapter, a model for the elastic behavior of a granular medium which allows the generation of grain contacts during hydrostatic compression is presented. The form of this model permits the use of a variety of contact microphysics and statistical distributions for the gap width of the near-contact. Numerical results for this model display significant differences in comparison with models that assume no contact generation, particularly in the transition region. Further, it can be observed that this transition zone will occur over the range of confining pressures of interest in geological and geophysical problems (10^6 - 10^7 Pa) for values of $\alpha$ small in comparison to the sphere radius. This means that extremely small near-contact gaps can have a significant effect on the elastic behavior of a granular medium. The behavior of the contact generating model is consistent with experimental results for random packings of glass beads.
\[ \lambda = 2.1474 \times 10^{10} \text{ Pa} \]
\[ \mu = 2.9655 \times 10^{10} \text{ Pa} \]
\[ R = 4.064 \times 10^{-5} \text{ m} \]
\[ \phi = 0.383 \]
\[ \eta_0 = 6.4 \]
\[ \eta_c = 2.1 \]

Table 2.1: Parameters used in numerical results.
Figure 2.1: Representative sphere in V (n th sphere), surrounded by neighboring spheres in contact (m th spheres) and near-contact (m' th spheres).
Figure 2.2: The original point contact configuration.
Figure 2.3: The near-contact configuration.
Figure 2.4: Computed log (-ε) as a function of log (p) using the contact generating model where (1) $\alpha = 0.001 R$ and (2) $\alpha = 0.0001 R$. Dashed lines are Walton (1987) model for $\eta = \eta_0$ (Line A- A') and $\eta = \eta_0 + \eta_C$ (Line B - B').
Figure 2.5: Computed log ($\lambda^*$) as a function of log (p) using the contact generating model where (1) $\alpha = 0.001$ R and (2) $\alpha = 0.0001$ R. Dashed lines are Walton (1987) model for $\eta = \eta_0$ (Line A- $A'$) and $\eta = \eta_0 + \eta_c$ (Line B- $B'$).
Figure 2.6: Computed log ($\mu^*$) as a function of log (p) using the contact generating model where (1) $\alpha = 0.001$ R and (2) $\alpha = 0.0001$ R. Dashed lines are Walton (1987) model for $\eta = \eta_o$ (Line A- A') and $\eta = \eta_o + \eta_c$ (Line B - B').
Chapter Three

Inclusion-Based Formulations for the Dielectric Properties of Heterogeneous Media

Introduction

In the range of frequencies of interest in the geophysical measurement of dielectric behavior (i.e. 100 kHz to 1 GHz), a dominant mechanism affecting the dielectric response of porous rocks is the charge polarization at interfaces on the microscopic scale (Poley et al., 1978). Inclusion-based formulations permit the incorporation of sufficient information about the geometrical configuration of the constituents in a porous rock to describe this phenomenon. Numerous studies have used inclusion-based formulations to investigate the relationship between microstructural features and macroscopic dielectric properties for heterogeneous media; these have been reviewed by van Beek (1967). All of these formulations describe the microstructure in terms of a background matrix into which inclusions are embedded; however, a diverse collection of mathematical expressions has been obtained. While some of this diversity can be attributed to the explicit use of specific inclusion types (i.e. inclusions with a given shape and composition), other variations result from fundamental differences in the methods used to derive these formulations. To illustrate these fundamental differences, as well as the common elements, the major classes of inclusion-based formulations are derived in this chapter. The degree of generalization that is maintained in this discussion is not found in previous work on the subject; therefore, the differences between the resulting expressions
are not caused by the explicit choice of specific inclusion types. In addition, this work has permitted us to identify a general class of iterative formulations and their associated hierarchical structure. It is important to note that the basic methods and principles used in the derivation of the dielectric properties are applicable to inclusion-based formulations for other physical properties.

**Definition of an Effective Medium**

To start the discussion, let us consider a heterogeneous medium that is composed of isotropic materials possessing causal, linear electrical rheologies that are described in terms of generalized permittivities (see Appendix 1). Let a volume \( V \) be some representative volume element within this heterogeneous system which has dimensions that are large in comparison to the scale of the inhomogeneities, yet small in comparison to the size of the total system. The material in \( V \) is subjected to an applied uniform electric field that is time harmonic with angular frequency \( \omega \). It will be assumed that the wavelengths and attenuation lengths of this field are large in comparison with the dimensions of \( V \). Therefore, a quasi-static approach is appropriate for determining the propagation of electromagnetic waves within this volume.

The microscopic structure is uniformly distributed such that the system is homogeneous and isotropic on a macroscopic scale. Hence, the macroscopic electrical behavior of the heterogeneous system can be simulated by the response of an effective homogeneous material with a generalized permittivity \( \varepsilon^*_g \). Let \( \mathbf{D} \) and \( \mathbf{E} \) be the generalized electric displacement and electric field strength within the system, respectively, in complex notation. Then, the effective generalized permittivity of the
material in $V$ is defined by

$$\overline{D} = \epsilon_g^* \overline{E},$$  \hspace{1cm} (3.1)$$

where

$$\overline{D} = V^{-1} \int_V D(x) \, dv$$  \hspace{1cm} (3.2)$$

and

$$\overline{E} = V^{-1} \int_V E(x) \, dv$$  \hspace{1cm} (3.3)$$

are the volume averages of $D$ and $E$, respectively. In addition, the pointwise constitutive relationship

$$D(x) = \epsilon_g(x) E(x)$$  \hspace{1cm} (3.4)$$

is valid at all points $x$ in $V$. Rearranging terms in Equation (3.4) and taking the volume average, one obtains

$$\overline{D} = \epsilon_g^* \overline{E} + V^{-1} \int_V \left[ \epsilon_g(x) - \epsilon_g^* \right] E(x) \, dv.$$  \hspace{1cm} (3.5)$$
Given the definition of the effective generalized permittivity $\varepsilon_g^*$ (Equation 3.1),

Equation (3.5) implies that

$$V^{-1} \int_V \left[ \varepsilon_g(x) - \varepsilon_g^* \right] E(x) \, dv = 0. \quad (3.6)$$

Hence, the volume average of the pointwise difference in polarization between the heterogeneous and effective media must vanish. This equation defines the fundamental relationship between a heterogeneous system and its equivalent effective medium. To evaluate Equation (3.6), the structure of the underlying medium used in inclusion-based formulations will now be considered; it will be seen that it is necessary to analyze the response of the individual inclusions in order to accomplish this procedure.

**Analysis of an Individual Inclusion**

It will be assumed that the microstructure within the heterogeneous medium can be described in terms of a homogeneous matrix with a generalized permittivity $\varepsilon_g^{(0)}$ into which homogeneous inclusions composed of materials having generalized permittivities $\varepsilon_g^{(i)}$ are embedded. Considering the structure of this system, Equation (3.6) may be rewritten as

$$V \varepsilon_g^* \overline{E} = V\varepsilon_g^{(0)} \overline{E} + \int_{V_{\text{inc}}} \left( \varepsilon_g^{(i)} - \varepsilon_g^{(0)} \right) E^{(i)}(x) \, dv \quad (3.7)$$

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where $V_{inc}$ is the volume of the inclusions in $V$ and $E^{(i)}(x)$ is the electric field strength within the inclusions. The two terms on the right hand side of Equation (3.7) can be viewed as a reference and perturbation term, respectively. In general, it is very difficult to precisely determine the electric field within the inclusions. However, it is possible to estimate $E^{(i)}(x)$ within an individual inclusion by treating it as a single inclusion embedded in an infinite, homogeneous background with a generalized permittivity $\varepsilon_g^{(b)}$. This system is subjected to an applied field with uniform strength $E_a$ at infinity. The values of $\varepsilon_g^{(b)}$ and $E_a$ are selected such that interactions with the other inclusions in the system are simulated. Given the macroscopic isotropy of the heterogeneous system, it will be assumed that $E_a$ and $\bar{E}$ are oriented in the same direction.

The quasi-static behavior of this system is govern by

$$\varepsilon_g^{(b)} \nabla^2 \psi^{(b)}(x) = 0$$  \hspace{1cm} (3.8)

within the background matrix and

$$\varepsilon_g^{(i)} \nabla^2 \psi^{(i)}(x) = 0$$  \hspace{1cm} (3.9)

inside the inclusion. The term $\psi$ is the electric potential; and, $E(x) = -\nabla \psi(x)$. Continuity of the electric potential and the normal component of the generalized electric displacement lead to the following boundary conditions along the surface $S$ that forms the interface between the inclusion and the background:
\[ \psi^{(b)}(x) \big|_S = \psi^{(i)}(x) \big|_S \] (3.10)

and

\[ \left[ (\varepsilon^{(b)}_g \nabla \psi^{(b)}(x)) \cdot \hat{n} \right] \big|_S = \left[ (\varepsilon^{(i)}_g \nabla \psi^{(i)}(x)) \cdot \hat{n} \right] \big|_S, \] (3.11)

where \( \hat{n} \) is the unit normal vector on \( S \) directed into the background. Such boundary conditions exclude the existence of surface phenomena, such as an electrical double layer, along the surface \( S \).

In order to obtain an analytic solution for \( E^{(i)}(x) \), the inclusion shapes must be restricted to ellipsoids. The ellipsoid under consideration has semi-axes of lengths \( a_1, a_2, \) and \( a_3 \), respectively, where \( a_1 \geq a_2 \geq a_3 \). The shape of the ellipsoid is uniquely defined by the aspect ratio pair \( (\alpha_1, \alpha_2) \), where

\[ 0 \leq \alpha_2 = \frac{a_3}{a_1} \leq \alpha_1 = \frac{a_2}{a_1} \leq 1. \] (3.12)

Using the results of Stratton (1941), it can be shown that the electric field within this ellipsoid is uniform and has a strength given by

\[ E^{(i)}(x) = E_a \varepsilon^{(b)}_g \sum_{j=1}^{3} \left[ \left( \varepsilon^{(b)}_g + (\varepsilon^{(i)}_g - \varepsilon^{(b)}_g) A_j(\alpha_1, \alpha_2) \right)^{-1} \cos \theta_j \hat{x}_j \right], \] (3.13)

where \( E_a \) is the magnitude of \( E_a \), \( \theta_j \) is the angle between the \( j \)th semi-axis and \( E_a \), and \( \hat{x}_j \) is the unit vector along the \( j \)th semi-axis. The term \( A_j(\alpha_1, \alpha_2) \) is the depolarization...
coefficient associated with the \( j \)th semi-axis and is uniquely determined by the shape of the ellipsoid. In terms of the semi-axis lengths, the depolarization coefficients are defined as

\[
A_j(\alpha_1, \alpha_2) = \frac{1}{2} a_1 a_2 a_3 \int_0^\infty \left( s + a_j^2 \right)^{-1} \left[ \prod_{k=1}^3 (s + a_k^2) \right]^{-\frac{1}{2}} ds \quad (j=1, 2, 3). \tag{3.14}
\]

The depolarization coefficients for the three principal axes are related by

\[
A_1(\alpha_1, \alpha_2) + A_2(\alpha_1, \alpha_2) + A_3(\alpha_1, \alpha_2) = 1. \tag{3.15}
\]

In order to preserve macroscopic isotropy in the heterogeneous medium, identical ellipsoidal inclusions are assumed to be oriented in a uniformly random manner. The electric field strength inside such an inclusion is estimated by averaging the result given by Equation (3.13) over all possible orientations of the inclusion with respect to \( E_a \).

Performing this process, it is found that the expected value of the estimated electric field strength within an inclusion is

\[
\langle E^{(i)}(x) \rangle = \frac{\varepsilon^{(b)}_g}{3} \left\{ \sum_{j=1}^3 \left[ \varepsilon^{(b)}_g + \left( \varepsilon^{(i)}_g - \varepsilon^{(b)}_g \right) A_j(\alpha_1, \alpha_2) \right]^{-1} \right\} E_a. \tag{3.16}
\]

This quantity will be used in Equation (3.7) to approximate the electric field inside the inclusions within the heterogeneous medium. It will now be shown that major classes of inclusion-based formulations can be defined by the manner in which \( \varepsilon^{(b)}_g \) and \( E_a \) are
used to specify the inclusion interactions. Further, no explicit restrictions will be placed on the shape and composition of the ellipsoidal inclusions permitted in the medium; thus, the expressions which are obtained are of a more general nature than those previously given in the literature.

Classical Inclusion Formulations

Let us consider a heterogeneous medium in which a continuous distribution of ellipsoidal inclusion types are present. It will be assumed that the nature of the inclusion interactions is such that the same values of \( \varepsilon_g^{(b)} \) and \( E_a \) can be used in the determination of \( E^{(i)}(\mathbf{x}) \) within the individual inclusions. Using the result given by Equation (3.16), Equation (3.7) becomes

\[
\varepsilon_g^* \bar{E} = \varepsilon_g^{(0)} \bar{E} + \frac{1}{3} \varepsilon_g^{(b)} E_a \int \int \int_{\varepsilon_g^{(i)}} c_{\text{inc}}(\varepsilon_g^{(i)}, \alpha_1, \alpha_2) \times
\]

\[
f(\varepsilon_g^{(b)}, \varepsilon_g^{(0)}, \varepsilon_g^{(i)}, \alpha_1, \alpha_2) d\alpha_2 d\alpha_1 d\varepsilon_g^{(i)}, \quad (3.17)
\]

where

\[
f(\varepsilon_g^{(b)}, \varepsilon_g^{(0)}, \varepsilon_g^{(i)}, \alpha_1, \alpha_2) = (\varepsilon_g^{(i)} - \varepsilon_g^{(0)}) \sum_{j=1}^{3} \left[ \varepsilon_g^{(b)} + (\varepsilon_g^{(i)} - \varepsilon_g^{(b)}) \Lambda_j(\alpha_1, \alpha_2) \right]^{-1} \quad (3.18)
\]

and \( \bar{E} \) is the magnitude of \( \bar{E} \). The distribution of inclusion types in \( V \) is described in
terms of the fractional volume density function \( c_{\text{inc}}(\varepsilon_g^{(i)}, \alpha_1, \alpha_2) \). For a distribution which is discrete in any inclusion parameter (i.e. \( \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \)), the appropriate integral is transformed into a summation and the definition of \( c_{\text{inc}} \) is modified.

Equation (3.17) represents a group of formulations which will be referred to as classical inclusion formulations; various subdivisions of this group are defined by the values chosen for \( \varepsilon_g^{(b)} \) and \( E_a \) in order to simulate inclusion interactions. The magnitude of these interactions is dependent on the inclusion types present and their concentrations, as well as the relative contrast in the generalized permittivity of the constituents. If the inclusion concentration is very dilute, reasonable accuracy is obtained by assuming that there are no inclusion interactions (i.e. \( \varepsilon_g^{(b)} = \varepsilon_g^{(0)} \) and \( E_a = \overline{E} \)) (Wagner, 1914; Burger, 1919; Fricke, 1953). Hence, Equation (3.17) can be expressed as

\[
\varepsilon_g^* = \varepsilon_g^{(0)} \left[ 1 + \frac{1}{3} \int_{\varepsilon_g^{(i)}} \int_{\alpha_1} \int_{\alpha_2} c_{\text{inc}} \left( \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) \times \int_{\varepsilon_g^{(0)}} f(\varepsilon_g^{(0)}, \varepsilon_g^{(0)}, \varepsilon_g^{(i)}, \alpha_1, \alpha_2) \, d\alpha_2 \, d\alpha_1 \, d\varepsilon_g^{(i)} \right],
\]

which is a generalization of what has previously been referred to as the dilute approximation.

At higher levels of inclusion concentration, it is necessary to incorporate inclusion interactions into the analysis. One approach for simulating this effect is to assume \( \varepsilon_g^{(b)} = \varepsilon_g^{(0)} \) and determine how the value of \( E_a \) is modified by increasing inclusion concentration. One method, which has its origins in the work of Lorentz and others...
(Mossotti, 1850; Clausius, 1879; Lorentz, 1880; Lorenz, 1880), achieves this by equating the sum of the induced dipole moments due to the inclusions in a given spherical volume of the heterogeneous medium with the induced dipole moment of an equivalent sphere composed of the effective medium (Landauer, 1978). From this analysis, it can be shown that

$$E_a = \frac{\varepsilon_g^* + 2\varepsilon_g^{(0)}}{3\varepsilon_g^{(0)}} \bar{E}.$$  \hspace{1cm} (3.20)

Hence, equation (3.17) can be expressed as

$$\frac{\varepsilon_g^* - \varepsilon_g^{(0)}}{\varepsilon_g^* + 2\varepsilon_g^{(0)}} = \frac{1}{9} \int_{\varepsilon_g^{(i)}} \int_{\alpha_1} \int_{\alpha_2} c_{inc} \left( \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) \times$$

$$f \left( \varepsilon_g^{(0)}, \varepsilon_g^{(0)}, \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) d\alpha_2 \, d\alpha_1 \, d\varepsilon_g^{(i)},$$  \hspace{1cm} (3.21)

which is a generalization of what is commonly referred to as the Lorentz approximation.

In addition, several other techniques for estimating the apparent field intensity $E_a$ have been suggested. Fricke (1924,1953) presented an approach where $E_a$ is assumed to be equal to the average electric field strength in the matrix material. Sihvola and Kong (1988) gave a technique where $E_a$ acting on a given inclusion is the sum of $\bar{E}$ and the strength of the depolarization fields due to the other inclusions.

An alternate approach for simulating inclusion interactions is to assume that $E_a = \bar{E}$.
and determine the appropriate value of $\varepsilon_g^{(b)}$ that describes the effective background experienced by the inclusions. From an analysis of experimental results, de Loor (1953, 1964, 1968) determined that $\varepsilon_g^{(b)}$ has a value between $\varepsilon_g^*$ and $\varepsilon_g^{(0)}$ in lossless composites and that no unique value of $\varepsilon_g^{(b)}$ adequately describes all heterogeneous media. Of particular interest is the case where $\varepsilon_g^{(b)} = \varepsilon_g^*$ (Böttcher, 1945; Polder and van Santan, 1946; Taylor, 1965), which is commonly referred to as the self-consistency condition. This value of $\varepsilon_g^{(b)}$ was determined by Böttcher (1945) on the basis of the theory of reaction fields (Onsager, 1936; Böttcher, 1942). Incorporating this condition into Equation (3.17), one obtains

$$
\varepsilon_g^* = \varepsilon_g^{(0)} + \frac{\varepsilon_g^*}{3} \int_{\varepsilon_g^{(1)}} \int_{\alpha_1} \int_{\alpha_2} c_{inc} \left( \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) \times
$$

$$
f \left( \varepsilon_g^*, \varepsilon_g^{(0)}, \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) d\alpha_2 d\alpha_1 d\varepsilon_g^{(i)}, \quad (3.22)
$$

which is a generalization of the average field or self-consistent approximation.

All classical inclusion formulations are obtained by determining the perturbation term in Equation (3.7) with respect to the background matrix $\varepsilon_g^{(0)}$. Hence, the accuracy of this group of formulations can be limited in terms of inclusion concentration. For cases of high levels of inclusion concentration, another group of formulations based on the two methods proposed by Bruggeman (1935) is commonly used.
Bruggeman's Symmetric and Unsymmetric Formulations

Bruggeman's unsymmetric approach, commonly referred to as the differential effective medium or iterated dilute approximation (Norris et al., 1985a; Avellaneda, 1987), utilizes an iterative construction process to fabricate the effective medium. Starting with a homogeneous background matrix with generalized permittivity \( \varepsilon_g^{(i)} \), each iteration involves the embedding of an infinitesimal portion of the inclusions present in the heterogeneous system. After each embedding, the dilute approximation (Equation 3.19) is used to determine the effective properties of the resulting medium; this resulting medium is used as the effective background matrix in the next embedding step.

Using the variable volume process of Norris et al. (1985a), the generalized form of the iterated dilute approximation can be derived as follows. Let \( V_0 \) be the initial volume of the system prior to the homogenization process; then, the volume of the system after the homogenization process \( (V_f) \) is given by

\[
V_f = V_0 \left( 1 - \Phi_{\text{inc}} \right)^{-1}
\]

(3.23)

where

\[
\Phi_{\text{inc}} = \int_0^1 \left[ \int_{\varepsilon_g^{(i)}} \int_{\alpha_1} \int_{\alpha_2} c_{\text{inc}} \left( t; \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) d\alpha_2 d\alpha_1 d\varepsilon_g^{(i)} \right] dt
\]

(3.24)

is the total fractional volume of the inclusions. The fractional volume density function \( c_{\text{inc}} \left( t; \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) \) is expressed in term of a homogenization parameter \( t \) that specifies the
order in which inclusions are incorporated into the effective medium during the homogenization process. For brevity, the term within the brackets in Equation (3.24) will be denoted by $\tilde{c}_{\text{inc}}(t)$. Therefore, the total volume of the heterogeneous medium at any point in the homogenization process, as specified by the value of $t$, is given by

\[
V(t) = (1 - \phi_{\text{inc}})^{-1} V_0 \left[ 1 - \int_t^1 \tilde{c}_{\text{inc}}(s) \, ds \right]. \tag{3.25}
\]

Let us now consider the change in the generalized effective permittivity of the system due to the embedding of a very small volume of inclusions which denotes the increment from $t$ to $t + \Delta t$ in the homogenization parameter. If $\varepsilon_g^*(t)$ and $\varepsilon_g^*(t + \Delta t)$ denote the general permittivities of the effective background matrix and resulting medium at this embedding step, then the dilute approximation (Equation 3.19) gives

\[
V(t + \Delta t) \varepsilon_g^*(t + \Delta t) = V(t + \Delta t) \varepsilon_g^*(t) + \frac{1}{3} \varepsilon_g^*(t) \times \nonumber
\]

\[
\left\{ V_0 (1 - \phi_{\text{inc}})^{-1} \int_t^{t + \Delta t} \left[ \int_{E_g} \int_{\alpha_1} \int_{\alpha_2} c_{\text{inc}}(s; \varepsilon_g^{(i)}, \alpha_1, \alpha_2) \times \right. \nonumber
\]

\[
f(\varepsilon_g^*(s), \varepsilon_g^*(s), \varepsilon_g^{(i)}, \alpha_1, \alpha_2) \, d\alpha_2 \, d\alpha_1 \, d\varepsilon_g^{(i)} \left] \, ds \right\} . \tag{3.26}
\]

As the volume of the embedded inclusions becomes infinitesimally small (i.e. $\Delta t \to 0$), then the change in $\varepsilon_g^*$ due to the homogeneous process is described by the differential
The generalized permittivity of the heterogeneous medium is determined by integrating Equation (3.27) from \( t = 0 \) to \( t = 1 \) with the initial condition \( \varepsilon_g^*(t = 0) = \varepsilon_g^{(0)} \). Variations in the embedding sequence can result in significantly different values of \( \varepsilon_g^* \) (Norris et al, 1985a). This initially appears to lead to ambiguities in the application of the iterative dilute approximation. However, it will be shown later that the nature of the microstructure within the heterogeneous medium uniquely defines this homogenization process. In addition, it should be noted that the embedding sequence determines the nature of effective background matrix for a particular inclusion as the value of \( \varepsilon_g^{(b)} \) varies between \( \varepsilon_g^{(0)} \) and \( \varepsilon_g^* \) during the homogenization process.

Bruggeman's symmetric effective medium theory, which is also referred to as the coherent potential or effective medium approximation (Landauer, 1952; Stroud, 1975; Milton, 1985), carries inclusion-based formulations to their logical extreme by considering the heterogeneous medium to be an aggregate of inclusions. This structure implies that Equation (3.6) can be rewritten as
\[
\int_V \left( \varepsilon_g^{(i)} - \varepsilon_g^* \right) E^{(i)}(x) \, dv = 0 \quad (3.28)
\]

Assuming \( E_g = E \) and the self-consistency condition, one obtains

\[
\int_{\varepsilon_g^{(i)}} \int_{\alpha_1} \int_{\alpha_2} c_{\text{inc}} \left( \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) f\left( \varepsilon_g^*, \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) d\alpha_2 \, d\alpha_1 \, d\varepsilon_g^{(i)} = 0 \quad (3.29)
\]

where

\[
\int_{\varepsilon_g^{(i)}} \int_{\alpha_1} \int_{\alpha_2} c_{\text{inc}} \left( \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) d\alpha_2 \, d\alpha_1 \, d\varepsilon_g^{(i)} = 1 \quad (3.30)
\]

Norris et al. (1985 a, b) and Avellaneda (1987) have established relationships between Bruggeman’s symmetric and unsymmetric approaches. In particular, these results imply that the value of \( \varepsilon_g^* \) obtained from the iterated dilute approximation as \( \phi_{\text{inc}} \to 1 \) for the case when all embedding steps are identical (i.e. \( c_{\text{inc}} \left( t, \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) = c_{\text{inc}} \left( \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) \)) converges to the \( \varepsilon_g^* \) given by the coherent potential approximation. Further, Milton (1984, 1985) has shown that the coherent potential approximation is physically realized by a homogenization-type process. This suggests that a generalized version of the effective medium approximation can be obtained by permitting \( c_{\text{inc}} \) to vary during the embedding process while maintaining \( \phi_{\text{inc}} = 1 \).

At this point, the major classes of inclusion-based formulations have been derived in their generalized form. The question that now arises is which of these expressions is the
most appropriate for a given heterogeneous medium. One criterion is to compare the
microstructure of the heterogeneous medium with the underlying microstructural model
for a given class of theoretical formulations.

**Implicit Microstructural Relationships**

Each theoretical formulation has an underlying microstructural model that determines
the applicability of the formulation to a given heterogeneous medium. While some
features of the microstructure are explicitly specified (e.g. designation of a background
phase and specification of inclusion types), other relationships are implicit in nature. Of
particular interest are the implicit microstructural relationships which result from use of
the Bruggeman techniques for higher levels of inclusion concentration.

The formulations derived in the preceding sections exhibit no explicit dependence of
the dielectric response on inclusion size. In fact, it has been proven that the electrical
properties are invariant with respect to the size of the inhomogeneities in the system
under the conditions assumed in this study (Cohen, 1981). However, an analysis of
experimental data for a number of simple systems indicates that there is a relationship
between the distribution of the inclusion sizes present within the system and the accuracy
of a given theoretical formulation (Pearce, 1955; De La Rue and Tobias, 1959; Meredith

It can be argued heuristically that individual inclusions of a given size are embedded
in an effective background material composed of the original background matrix material
and the smaller size inclusions. Therefore, an iterative construction process will be used
to determine the effective generalized permittivity of the system. Starting with an
original background matrix $\varepsilon_g^{(0)}$, inclusions are sequentially embedded in order of
increasing size. At the $k$ th iteration, all inclusions of a given size are embedded and the
effective generalized permittivity of the resultant system is determined; this system
becomes the effective background matrix for the $(k+1)$ th embedding step (i.e. $\varepsilon_g^* (k) =
\varepsilon_g^{(b)} (k + 1)$).

Let us initially consider a system containing $K$ discrete sizes of inclusions. The
variable volume process (Norris et al., 1985a) will be used during the iterative embedding
procedure. Using a classical inclusion-based formulation, the embedding of the $k$ th size
inclusions is described by

$$
\varepsilon_g^*(k) E = \varepsilon_g^{(b)} (k) E + \frac{\varepsilon_g^{(b)} (k) E_a}{3} \left[ 1 - \sum_{j=k+1}^{K} \int_{\varepsilon_g^{(0)}}^{\varepsilon_g^{(b)}(k)} \int_{\alpha_1}^{\alpha_2} c_{\text{inc}} \left( j; \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) d\alpha_2 d\alpha_1 d\varepsilon_g^{(i)} \right]^{-1} \times
\int_{\varepsilon_g^{(i)}}^{\varepsilon_g^{(b)}(k)} \int_{\alpha_1}^{\alpha_2} c_{\text{inc}} \left( k; \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) f \left( \varepsilon_g^{(b)}(k), \varepsilon_g^{(b)}(k), \varepsilon_g^{(b)}(k), \alpha_1, \alpha_2 \right) d\alpha_2 d\alpha_1 d\varepsilon_g^{(i)} \end{equation}$$

(3.31)

where $c_{\text{inc}}(k; \varepsilon_g^{(i)}, \alpha_1, \alpha_2)$ is the volume fraction distribution function of inclusions of the
$k$ th size and $E_a$ is determined by the approximation used. The effective generalized
permittivity of the heterogeneous medium is given by $\varepsilon_g^* = \varepsilon_g^*(K)$.

Let us consider the case when a piecewise continuous spectrum of inclusion sizes is
present in the system. This implies that the construction process involves the
incorporation of an infinitesimally small inclusion volume during each embedding. In
this case, the construction process described by Equation (3.31) becomes
\[
\frac{d\varepsilon_g^*(s)}{ds} = \varepsilon_g^*(s) \left( 1 - \int_s^{s_{\text{max}}} \left[ \int_{\varepsilon_g^{(i)}} \int_{\alpha_1} \int_{\alpha_2} c_{\text{inc}} \left( t; \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) dt \right] d\alpha_2 d\alpha_1 d\varepsilon_g^{(i)} \right)^{-1} x
\]

\[
\left[ \int_{\varepsilon_g^{(i)}} \int_{\alpha_1} \int_{\alpha_2} c_{\text{inc}} \left( s; \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) f\left( \varepsilon_g^*(s), \varepsilon_g^*(s); \varepsilon_g^{(i)}, \alpha_1, \alpha_2 \right) d\alpha_2 d\alpha_1 d\varepsilon_g^{(i)} \right]
\]

(3.32)

where \(s_{\text{min}}\) and \(s_{\text{max}}\) are the minimum and maximum inclusion sizes present in the medium. The effective generalized permittivity of the heterogeneous medium is obtained by integrating this initial value problem from \(\varepsilon_g^*(s_{\text{min}}) = \varepsilon_g^{(0)}\) to \(\varepsilon_g^*(s_{\text{max}}) = \varepsilon_g^*\).

Comparison of this result with Equation (3.27) shows the equivalence between this type of hierarchical medium and the iterated dilute approximation. Further, this equivalence demonstrates that the homogenization process in the iterated dilute approximation is uniquely defined in terms of the microstructure of the medium (i.e. the embedding sequence of the inclusions proceeds in order of progressively larger size).

However, the heuristic argument used to motivate the hierarchical models described by Equation (3.31) implicitly assumes that the inclusion of a given size are significantly larger than those incorporated in the effective background matrix. This implies that the relative concentration of inclusions having similar sizes is sufficiently dilute such that their mutual interactions can be neglected. For high levels of inclusion concentration, this implies that a very broad range of sizes must be present. This conclusion is supported by experimental observation (De La Rue and Tobias, 1959).

The nature of the self-consistency condition used in the average field and coherent potential approximations now becomes apparent. For all inclusions to experience the
effective medium $\varepsilon^*_g$ as a background matrix, the distribution of inclusion types must be identical for all sizes (i.e. $c_{\text{inc}}(s; \varepsilon^{(i)}_g, \alpha_1, \alpha_2) = c_{\text{inc}}(\varepsilon^{(i)}_g, \alpha_1, \alpha_2)$). This implies that the inclusions must be present in a wide range of sizes and fill the entire space (i.e. $\Phi_{\text{inc}} = 1$). The generalized version suggested for the effective medium approximation differs from the coherent potential approximation in that distribution of inclusion types varies with size; hence, this generalization does not utilize the self-consistency condition.

Another implicit relationship of the Bruggeman formulations is the degree of connectivity a given constituent possesses in the underlying model. To investigate this relationship for the coherent potential and iterated dilute approximations, Yonezawa and Cohen (1983) and Norris et al. (1985 a,b) employed simple two-component systems consisting of conducting and insulating phases to determine the percolation threshold of the resulting medium. In the iterated dilute approximation, the designated background component remains connected over all levels of inclusion concentration for non-degenerate cases of inclusion shape (i.e. needles and flat disc). In contrast, the coherent potential approximation always exhibits non-zero percolation thresholds beyond which complete connectivity of a given component ceases to exist; the value of this threshold is a function of the inclusion shapes.

Conclusions

The generalized versions of the major classes of inclusion-based formulations used to determine the macroscopic physical properties of heterogeneous media have been derived in this chapter. All formulations of this type originate from the requirement that the pointwise deviation in behavior between a heterogeneous system and its equivalent
effective medium must vanish in some average sense. To evaluate this criterion, it is necessary to determine the responses of the individual inclusions; their behavior is estimated by analyzing single inclusions embedded in an infinite, homogeneous background. Inclusion interactions are simulated by means of the imposed conditions used in this analysis; the major classes of inclusion-based formulations can be defined on the basis of the technique used to simulate these interactions. Each class of inclusion-based formulations has an associated microstructure that specifies the topological relationships between the various components. Of particular significance are the hierarchical structures, in terms of inclusion sizes, that are implied by use of an iterated embedding process. Further, the physical realization of formulations employing the self-consistency condition represent a specific case of these hierarchical structures. In the next two chapters, inclusion-based formulations will be used to investigate the effects of variations in microstructure due to changes in pore-scale fluid distribution on the dielectric and elastic responses of partially-saturated porous rocks.
Chapter Four

The Effects of Pore-Scale Fluid Distribution on Elastic and Electromagnetic Wave Propagation in Partially Saturated Porous Rocks: A Theoretical Analysis

Introduction

It is well established that the elastic and electrical properties of rocks vary as a function of the level of water saturation ($S_w$) in the pore space. Hence, it is assumed that the level of water saturation within a porous rock can be determined by measuring these properties. However, it is now apparent that the specific form of the functional relationship between a particular physical property and $S_w$ is significantly affected by the pore-scale fluid distribution occurring within the medium. This has been observed in laboratory data for elastic wave velocities (Domenico, 1976, 1977; Knight and Nolen-Hoeksema, 1990), elastic wave attenuation (Bourbie and Zinszner, 1984), electrical resistivity (Longeron et al., 1989), and dielectric constant (Knight and Nur, 1987a). Significantly different values of a given physical property were observed for the same level of water saturation in these studies by varying the pore-scale fluid distribution through the use of differing saturation methods.

Many of the theoretical formulations commonly used to model the physical properties of partially saturated rocks cannot account for this variation in functional dependence with changes in pore-scale fluid distribution. This is due to the inadequate description of the geometrical configuration of the constituents within the medium. In this chapter,
mathematical formulations are derived that allow the incorporation of the necessary geometrical information about pore-scale fluid distribution. These formulations are used to undertake a theoretical analysis of the effects of pore-scale fluid distribution on elastic and electromagnetic wave propagation in partially saturated porous rocks.

Geometrical Description of Porous Rocks

Let us consider the description of the geometrical configuration of the constituents within a porous rock. In order to utilize the inclusion-based formulations discussed in Chapter Three, this system will be viewed as being composed of a rock matrix into which inclusions representing individual pore spaces are embedded. By using this discretization of the system, it is possible to specify the saturation conditions within the individual pore spaces; however, it also implies that intra-pore phenomena control the physical properties of interest. For elastic and electromagnetic wave propagation, communication between individual pores is minimized when sampling frequencies are sufficiently high and/or the permeability of the pore system is adequately low. Under these conditions, pore connectivity can be neglected in the microstructural description of a porous rock when determining its dielectric and elastic properties.

Since the volume fraction of the pore space is usually small, accurate results are commonly obtained by using the classical inclusion formulations with this form of discretization. In order to avoid possible concentration limitation associated with the dilute approximation, the Lorentz approximation will be used to compute the dielectric and elastic response of the porous rock system. The differences between results obtained from use of the Lorentz approximation and the higher order iterative dilute
approximation are small for parameters that realistically describe porous rocks.

Information from pore cast (Swanson, 1979) and SEM observations (Timur et al., 1971; Brace et al., 1972; Sprunt and Brace, 1974) indicates that the pore space geometries can be adequately described in terms of oblate spheroids. The shape of an oblate spheroid is uniquely determined by a single aspect ratio \( \alpha \) (i.e. the ratio of the minor semi-axis length to the major semi-axis length) that has a value between 0 and 1, inclusively. The description of the pore space geometries used in this analysis will be restricted to these shapes. Hence, the pore shapes used in this description range from spheres representing the major equidimensional pore volumes to discs representing cracks.

By adopting this restriction on the permissible pore shapes, the computational effort required to determine the physical properties is significantly reduced. In particular, the depolarization coefficients which appear in the inclusion-based formulations for the dielectric response can be expressed in a simple analytical form. For a non-degenerate oblate spheroid (i.e. \( 0 < \alpha < 1 \)), the depolarization coefficients are given by (Landau and Lifshitz, 1960)

\[
A_3 (\alpha_1, \alpha_2) = A (\alpha) = \frac{1 + e^2}{e^3} (e - \tan^{-1} e) \tag{4.1}
\]

and

\[
A_1 (\alpha_1, \alpha_2) = A_2 (\alpha_1, \alpha_2) = \frac{1}{2} (1 - A (\alpha)) \tag{4.2}
\]
where

\[ e = (\alpha^2 - 1)^{1/2}. \] (4.3)

For the degenerate cases of a sphere (i.e. \( \alpha=1 \)) and a flat disc (i.e. \( \alpha=0 \)), the depolarization coefficient associated with the unique principal axis is \( \frac{1}{3} \) and 1, respectively.

The shapes of the individual pore spaces are assumed to be independent of changes in the local saturation conditions. Hence, the distribution of pore geometries within a given rock is considered to be invariant and will be described in terms of a pore geometry spectrum \( c_{\text{pore}}(\alpha) \). While \( c_{\text{pore}}(\alpha) \) can be viewed as a continuous spectrum, a discrete spectrum consisting of \( M \) differently shaped oblate spheroids will be used to describe the pore space geometries.

Using the above description of the porous rock system, the mathematical expressions for the dielectric and elastic response of partially saturated porous rocks will now be determined. The forms of these expressions depend on the nature of the pore fluid configuration within the individual pore spaces. While porous rock can contain numerous pore fluids, the systems considered in this study will be assumed to contain only two distinct pore fluids. These systems are sufficient to illustrate the influence of pore-scale fluid distribution on the macroscopic physical properties.

**Mathematical Formulations for Homogeneous Pore Spaces**

Let us initially consider a situation where the individual pore spaces are completely
filled with one of the two pore fluids; this implies that they are represented by homogeneous inclusions. Using Equation (3.21), the Lorentz approximation for the dielectric response of this system is given by

$$\frac{\varepsilon_g^* - \varepsilon_g^{(0)}}{\varepsilon_g^* + 2\varepsilon_g^{(0)}} = \frac{1}{9} \left[ \sum_{n=1}^{2} \sum_{m=1}^{M} s_m^{(n)} c_{\text{pore}}(\alpha_m) g(\varepsilon_g^{(0)}, \varepsilon_g^{(n)}, \alpha_m) \right]$$

(4.4)

where

$$g(\varepsilon_g^{(0)}, \varepsilon_g^{(n)}, \alpha_m) = \left( \frac{\varepsilon_g^{(n)} - \varepsilon_g^{(0)}}{\varepsilon_g^{(n)} + \varepsilon_g^{(0)}} \right) \left\{ \left[ \frac{\varepsilon_g^{(0)} + (\varepsilon_g^{(n)} - \varepsilon_g^{(0)})}{4} \right] A(\alpha_m) \right\}^{-1} +$$

$$4\left[ \frac{\varepsilon_g^{(0)} + \varepsilon_g^{(n)}}{4} \right] A(\alpha_m) \right\}^{-1} .$$

(4.5)

In this expression, $\varepsilon_g^{(0)}$ is the generalized permittivity of rock matrix; and, $\varepsilon_g^{(n)}$ (n=1, 2) are the generalized permittivity of the pore fluids. Further, the saturation level of the n th pore fluid in the m th shaped pores is given by $s_m^{(n)}$ where

$$s_m^{(1)} + s_m^{(2)} = 1 \ (m=1, \ldots, M).$$

(4.6)

Analogous expressions for the effective elastic moduli of this system can be obtained from the formulation derived by Kuster and Toksöz (1974). The effective bulk modulus
\( \frac{k^* - k^{(0)}}{3k^* + 4\mu^{(0)}} = \frac{1}{3} \left[ \sum_{n=1}^{2} \sum_{m=1}^{M} s_m^{(n)} c_{\text{pore}}(\alpha_m) \left( \frac{k^{(n)} - k^{(0)}}{3k^{(0)} + 4\mu^{(0)}} \right) T_{ijij}(k^{(0)}, \mu^{(0)}, k^{(n)}, \mu^{(n)}, \alpha_m) \right] \) (4.7)

and

\[ \frac{\mu^* - \mu^{(0)}}{6\mu^* (k^{(0)} + 2\mu^{(0)}) + \mu^{(0)} (9k^{(0)} + 8\mu^{(0)})} = \frac{1}{25} \left\{ \sum_{n=1}^{2} \sum_{m=1}^{M} s_m^{(n)} c_{\text{pore}}(\alpha_m) \times \right. \]

\[ \left. \frac{\mu^{(1)} - \mu^{(0)}}{\mu^{(0)} (3k^{(0)} + 4\mu^{(0)})} \left[ T_{ijij} - \frac{1}{3} T_{iiij} \right](k^{(0)}, \mu^{(0)}, k^{(n)}, \mu^{(n)}, \alpha_m) \right\}, \] (4.8)

respectively. The elastic moduli of the rock matrix are denoted by \( k^{(0)}, \mu^{(0)} \) and those of the pore fluids are represented by \( k^{(n)}, \mu^{(n)} \) (n=1, 2). The functions \( T_{ijij} \) and \( \left[ T_{ijij} - \frac{1}{3} T_{iiij} \right] \) are complicated expressions involving \( \alpha \) and the constituent elastic moduli; the reader is referred to Kuster and Toksoz (1974) where both functions are defined. The effective density \( \rho^* \) is given by

\[ \rho^* = \left( 1 - \phi_{\text{pore}} \right) \rho^{(0)} + \sum_{n=1}^{2} \sum_{m=1}^{M} s_m^{(n)} c_{\text{pore}}(\alpha_m) \rho^{(n)} \] (4.9)

where

\[ \phi_{\text{pore}} = \sum_{m=1}^{M} c_{\text{pore}}(\alpha_m) \] (4.10)
is the total porosity in volume fraction, \( \rho^{(0)} \) is the density of the rock matrix, and \( \rho^{(n)} \) (n=1, 2) are the densities of the pore fluids. Once the effective elastic moduli and density are determined, the effective compressional and shear wave velocities are obtained from

\[
V_p^* = \left( \frac{k^* + \frac{4}{3} \mu^*}{\rho^*} \right)^{\frac{1}{2}}
\]

(4.11)

and

\[
V_s^* = \left( \frac{\mu^*}{\rho^*} \right)^{\frac{1}{2}}
\]

(4.12)

respectively.

To this point, only pore spaces that are completely filled with a single pore fluid have been considered. However, it is probable that several distinct pore fluid species will simultaneously occupy the same pore space within a partially saturated porous rock. Let us now consider the dielectric and elastic response of a porous rock when this situation occurs.

**Mathematical Formulations for Heterogeneous Pore Spaces**

When several immiscible pore fluids coexist within an individual pore space, it is necessary to use a heterogeneous inclusion, i.e. an inclusion containing two or more constituents, in order to describe the configuration of the pore fluids. In the case of the
elastic properties, the contents of a heterogeneous pore space can be replaced in the theoretical formulations by a homogeneous inclusion filled with an effective pore fluid regardless of the geometrical configuration of the fluids within that pore space. To show this, let us consider an individual pore space with a total volume \( V \) occupied with volumes \( V_1 \) and \( V_2 \) of the inviscid fluid species 1 and 2, respectively. Only the hydrostatic pressure component of the stress tensor exists in either of the fluids.

Assuming that surface tension can be neglected, continuity of tractions at an interface between the fluids implies that \( p_1 = p_2 \) (the pressure in the respective pore fluid species).

Let \( \delta p \) be an incremental change in the hydrostatic pressure of the pore fluids, then

\[
\delta p = k^{(1)} \frac{\delta V_1}{V_1}
\]

(4.13)

and

\[
\delta p = k^{(2)} \frac{\delta V_2}{V_2},
\]

(4.14)

where \( \delta V_1 \) and \( \delta V_2 \) are the incremental change in the volumes occupied by fluids 1 and 2, respectively. The contents of the pore volume are replaced with an effective fluid with bulk modulus \( k^{(e)} \) and shear modulus \( \mu^{(e)} = 0 \), then

\[
\delta p = k^{(e)} \frac{\delta V}{V} = k^{(e)} \left( \frac{\delta V_1 + \delta V_2}{V} \right).
\]

(4.15)

Rearranging the terms in Equation (4.15) and using Equations (4.13) and (4.14), we
obtain

\[(k^{(e)} s^{(1)}))^{-1} = s^{(1)}(k^{(1)}))^{-1} + (1-s^{(1)})(k^{(2)}))^{-1}, \tag{4.16}\]

where \(s^{(1)}\) is the volume fraction of pore fluid 1 in the pore space. It should be noted that given the above assumptions, Equation (4.16) is valid for any configuration of pore fluids within the pore space.

To determine the elastic properties of the medium, the heterogeneous inclusions are replaced by homogeneous ones filled with the appropriate effective pore fluid. While a continuous spectrum of \(s^{(1)}\) values can occur within the heterogeneous pores of a given shape, it will be assumed in this study that all heterogeneous pores of a given shape have the same saturation level \(s^{(1)}_m\). Hence, the effective elastic moduli are given by

\[
\frac{k^* - k^{(0)}}{3 k^* + 4 \mu^{(0)}} = \frac{1}{3} \sum_{m=1}^{M} c_{\text{pore}}(\alpha_m) \left( \frac{k^{(e)}(s^{(1)}_m) - k^{(0)}}{3k^{(0)} - 4\mu^{(0)}} \right) \]

\[
T_{ij}(k^{(0)}, \mu^{(0)}, k^{(e)}(s^{(1)}_m), \mu^{(e)}(s^{(1)}_m), \alpha_m) \right) \tag{4.17}
\]

and

\[
\frac{\mu^* - \mu^{(0)}}{6 \mu^* (k^{(0)} + 2 \mu^{(0)} + \mu^{(0)} (9k^{(0)} + 8 \mu^{(0)})} = \frac{1}{25} \sum_{m=1}^{M} c_{\text{pore}}(\alpha_m) \times 
\]

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\[
\begin{align*}
\frac{\mu^{(e)}(s^{(1)}) - \mu^{(0)}}{\mu^{(0)}(3k^{(0)} + 4\mu^{(0)})}\left[T_{ijij} - \frac{1}{3} T_{iiij}\right](k^{(0)}, \mu^{(0)}, k^{(e)}(s^{(1)}), \mu^{(e)}(s^{(1)}), \alpha_m) \right] .
\end{align*}
\] (4.18)

In contrast to the elastic behavior, it is necessary to specify the geometrical configuration of the pore fluids within the heterogeneous pore spaces when mathematical expressions for electromagnetic wave propagation are derived. Two configurations of pore fluids for which analytical descriptions of the response can be obtained will now be considered. In both cases, \(\varepsilon_{g}^{(1)}\) and \(\varepsilon_{g}^{(2)}\) will denote the generalized permittivity of the non-wetting and wetting pore fluids, respectively. First, let us consider the case where an effective pore fluid is generated by embedding small spherical inclusions of the non-wetting phase into a background of the wetting fluid such that the resulting assemblage resembles a gas bubble-water mixture. Using the Lorentz approximation, the generalized permittivity of the effective pore fluid \(\varepsilon_{g}^{(e)}\) is given by

\[
\frac{\varepsilon_{g}^{(e)}(s^{(1)}) - \varepsilon_{g}^{(2)}}{\varepsilon_{g}^{(e)}(s^{(1)}) + 2\varepsilon_{g}^{(2)}} = s^{(1)} \frac{\varepsilon_{g}^{(1)} - \varepsilon_{g}^{(2)}}{\varepsilon_{g}^{(1)} + 2\varepsilon_{g}^{(2)}} ,
\] (4.19)

where \(s^{(1)}\) is the fractional volume of the non-wetting phase in the pore space. Equation (4.19) is valid for all values of \(s^{(1)}\) between 0 and 1, inclusively, and implies that the individual spheres of the non-wetting phase remain discrete and non-overlapping. To determine the effective generalized permittivity of a system containing this type of inclusions, the heterogeneous inclusions are replaced by homogeneous pores with the appropriate value of \(\varepsilon_{g}^{(e)}(s^{(1)})\). Therefore, the dielectric response of such a porous...
medium is described by

\[
\frac{\varepsilon_g^* - \varepsilon_g^{(0)}}{\varepsilon_g^* + 2\varepsilon_g^{(0)}} = \frac{1}{9} \left[ \sum_{m=1}^{M} c_{\text{pore}}(\alpha_m) g\left(\varepsilon_g^{(0)}, \varepsilon_g^{(e)}(s_m^{(1)}), \alpha_m\right) \right].
\] (4.20)

For the second case, the geometrical configuration of pore fluids within the individual pore spaces is represented by a two component, confocally-layered spheroid. Let the outer spheroidal shell be composed of the wetting phase and the inner spheroid consist of the non-wetting fluid. Stepin (1965) derived an expression for the response of a medium containing identical confocally-layered ellipsoidal inclusions. Expanding this result, the following expression is obtained for a porous medium containing confocally-layered pore spaces:

\[
\frac{\varepsilon_g^* - \varepsilon_g^{(0)}}{\varepsilon_g^* + 2\varepsilon_g^{(0)}} = \frac{1}{9} \left[ \sum_{m=1}^{M} c_{\text{pore}}(\alpha_m) h\left(\varepsilon_g^{(0)}, \varepsilon_g^{(1)}, \varepsilon_g^{(2)}, s_m^{(1)}, \alpha_m\right) \right]
\] (4.21)

where

\[
h\left(\varepsilon_g^{(0)}, \varepsilon_g^{(1)}, \varepsilon_g^{(2)}, s_m^{(1)}, \alpha_m\right) = \]

\[
\frac{N_1(\varepsilon_g^{(0)}, \varepsilon_g^{(1)}, \varepsilon_g^{(2)}, s_m^{(1)}, \alpha_m)}{D_1(\varepsilon_g^{(0)}, \varepsilon_g^{(1)}, \varepsilon_g^{(2)}, s_m^{(1)}, \alpha_m)} + \frac{N_2(\varepsilon_g^{(0)}, \varepsilon_g^{(1)}, \varepsilon_g^{(2)}, s_m^{(1)}, \alpha_m)}{D_2(\varepsilon_g^{(0)}, \varepsilon_g^{(1)}, \varepsilon_g^{(2)}, s_m^{(1)}, \alpha_m)},
\] (4.22)

\[
N_1(\varepsilon_g^{(0)}, \varepsilon_g^{(1)}, \varepsilon_g^{(2)}, s_m^{(1)}, \alpha_m) = e_g^{(2)}(e_g^{(1)} - e_g^{(0)}) s_m^{(1)} (1 - s_m^{(1)}) (e_g^{(2)} - e_g^{(0)}) \times
\]
\[
\left[ B(s_m^{(1)}, \alpha_m) \left( \varepsilon_g^{(2)} - \varepsilon_g^{(1)} \right) - \varepsilon_g^{(2)} \right],
\]  
(4.23)

\[
D_1(\varepsilon_g^{(0)}, \varepsilon_g^{(1)}, \varepsilon_g^{(2)}, s_m^{(1)}, \alpha_m) = \left[ B(s_m^{(1)}, \alpha_m) \left( \varepsilon_g^{(2)} - \varepsilon_g^{(1)} \right) - \varepsilon_g^{(2)} \right] \times
\]

\[
\left[ A(\alpha_m) \left( \varepsilon_g^{(0)} - \varepsilon_g^{(2)} \right) - \varepsilon_g^{(0)} \right] + s_m^{(1)} \left[ 1 - A(\alpha_m) \right] \left( \varepsilon_g^{(2)} - \varepsilon_g^{(1)} \right) \left( \varepsilon_g^{(0)} - \varepsilon_g^{(2)} \right),
\]
(4.24)

\[
N_2(\varepsilon_g^{(0)}, \varepsilon_g^{(1)}, \varepsilon_g^{(2)}, s_1, \alpha) = 8\varepsilon_g^{(2)} \left( \varepsilon_g^{(1)} - \varepsilon_g^{(0)} \right) s_m^{(1)} - 4(1 - s_m^{(1)}) \left( \varepsilon_g^{(2)} - \varepsilon_g^{(0)} \right) \times
\]

\[
\left[ B(s_m^{(1)}, \alpha_m) \left( \varepsilon_g^{(1)} - \varepsilon_g^{(2)} \right) - \varepsilon_g^{(1)} + \varepsilon_g^{(2)} \right]
\]
(4.25)

and

\[
D_2(\varepsilon_g^{(0)}, \varepsilon_g^{(1)}, \varepsilon_g^{(2)}, s_m^{(1)}, \alpha_m) = \left[ B(s_m^{(1)}, \alpha_m) \left( \varepsilon_g^{(1)} - \varepsilon_g^{(2)} \right) - \varepsilon_g^{(1)} + \varepsilon_g^{(2)} \right] \times
\]

\[
\left[ A(\alpha_m) \left( \varepsilon_g^{(2)} - \varepsilon_g^{(0)} \right) - \varepsilon_g^{(2)} + \varepsilon_g^{(0)} \right] + s_m^{(1)} \left[ 1 - A(\alpha_m)^2 \right] \left( \varepsilon_g^{(2)} - \varepsilon_g^{(1)} \right) \left( \varepsilon_g^{(0)} - \varepsilon_g^{(2)} \right). \quad (4.26)
\]

The term \( B(s_m^{(1)}, \alpha_m) \) is the depolarization coefficient associated with the unique principal axis of the inner spheroid. Evaluation of \( B(s_m^{(1)}, \alpha_m) \) requires the determination of the aspect ratio of the inner spheroid which is a function of both \( s_m^{(1)} \) and \( \alpha_m \) (See Appendix 2). For both types of heterogeneous inclusions, the cases \( s_m^{(1)} = 0 \) and \( s_m^{(1)} = 1 \) reduce to the expressions for homogeneous inclusions composed of the wetting phase and non-wetting phase, respectively.

From the above results, it can be seen that the mathematical formulations for the
dielectric and elastic properties of a partially saturated porous rock differ for the various geometrical configurations of the pore fluids within the individual pore spaces. To appreciate the significance of these differences, it is necessary to simulate the physical behavior of a model system using these expressions. In the following analysis, this approach will be used to illustrate the factors that control the relationship between pore-scale fluid distribution and macroscopic physical properties.

**Numerical Simulations**

A partially saturated porous rock containing air and deionized water will be used in the following numerical simulations. In order to employ the preceding mathematical expressions, it is necessary to specify the constituent properties. In the case of electromagnetic wave propagation, it will be assumed that the rock matrix and pore fluids are simple, linear materials where

\[
\varepsilon_{a}^{(j)} = \varepsilon_{a}^{(j)} + \frac{\sigma_{a}^{(j)}}{i\omega} \quad (j=0, 1, 2);
\]  

(4.27)

\(\varepsilon_{a}^{(j)}\) and \(\sigma_{a}^{(j)}\) are the apparent permittivity and conductivity of the j th component, respectively. Both apparent quantities are taken to be independent of \(\omega\). This is a reasonable approximation for the dielectric behavior of commonly occurring pore fluids, and both \(\varepsilon_{a}^{(j)}\) and \(\sigma_{a}^{(j)}\) are easily obtained. However, due to complex physical and chemical processes taking place at the rock-water interface, determination of the rock matrix parameters requires special consideration. Knight and Endres (1990) showed that

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the adsorbed water in a water-wetted sandstone should be considered as part of the rock matrix and the parameters used to describe this system should reflect this wetted state. This is an important factor since the parameters describing a wetted and dry rock matrix can differ significantly, particularly in the case of sandstones with large values of surface area to volume ratio (S/V). By incorporating the adsorbed water into the rock matrix system, the \( S_w \) used in the theoretical models refers to the level of bulk water saturation \( (S_{bw}) \).

The parameters used to describe the dielectric properties of the constituents are given in Table (4.1). The wetted rock matrix values were obtained from an analysis of dielectric data for a sample of Berea sandstone. The frequency used in the modeling computations is 105 kHz. The parameters used for modeling the elastic properties are given in Table (4.2). It should be noted that the difference between dry and wetted matrix elastic parameters was observed by Toksöz et al. (1976); their value for wetted Berea sandstone is used in this analysis.

Initially, the effects of the pore fluid configuration within the individual pore spaces will be investigated. For each type of pore fluid configuration, a series of models containing a single pore geometry (i.e. all inclusions in a given model have the same value of \( \alpha \)) is generated; models within a series are varied by changing the aspect ratio of the pore space. For a given model, the physical property of interest is determined as a function of \( S_{bw} \); for presentation purposes, these results are normalized to the \( S_{bw} = 0.00 \) value of the quantity under consideration. It is assumed that the value of local \( S_{bw} \) (i.e. the value of \( S_{bw} \) for an individual pore space) for all pore spaces is identical to the overall value of \( S_{bw} \) when a heterogeneous inclusion type is utilized. The total porosity \( \phi_{pore} \) is
maintained at 0.01 in all models.

Let us first consider the relationship between pore fluid configuration and the dielectric response. Figures (4.1) and (4.2) give the computed values of the apparent dielectric constant (i.e. $\kappa_a^* = \varepsilon_a^*/\varepsilon_{\text{vacuum}}$) and the apparent conductivity $\sigma_a^*$, respectively, as a function of $S_{bw}$ for models employing homogeneous inclusions. It can be observed that both $\kappa_a^*$ and $\sigma_a^*$ display nearly linear variations with changing $S_{bw}$. Figures (4.3) and (4.4) show the computed values of $\kappa_a^*$ and $\sigma_a^*$, respectively, as a function of $S_{bw}$ for models utilizing heterogeneous inclusions filled with an effective fluid consisting of air bubbles in water. These results differ significantly from those employing homogeneous inclusions in that distinctly non-linear relationships are observed for both $\kappa_a^*$ and $\sigma_a^*$.

Figures (4.5) and (4.6) illustrate the computed values of $\kappa_a^*$ and $\sigma_a^*$, respectively, as a function of $S_{bw}$ for models using heterogenous inclusions consisting of an air inner spheroid and confocal water shell. Once again, these results differ from those obtained with homogeneous inclusions in that distinctly non-linear relationships are observed for both $\kappa_a^*$ and $\sigma_a^*$. Further, comparison of results obtained with the two types of heterogeneous inclusions show that the variation in the pore fluid configuration drastically alters the functional dependence on $S_{bw}$.

Another way to alter the pore fluid configuration within the heterogeneous inclusions is to reverse the roles of the wetting and non-wetting phases. The effects of wetting phase variation on the dielectric properties of partially saturated rock has been experimentally observed by Poley et al. (1978). In the above simulations, the water phase was taken to be the wetting phase. While air cannot realistically act as the wetting phase
in the pore fluid configurations used in this analysis, its dielectric properties do not
significantly differ from those of many hydrocarbons in comparison to the dielectric
properties of water. Therefore, it is useful to consider the effect of reversing the role of
the two pore fluids used in this study. Figures (4.7), (4.8) and (4.9), (4.10) show the
computed values of $\kappa_a^*$ and $\sigma_a^*$, respectively, for the two types of heterogeneous inclusions
when air-wetted. Comparison of the air- and water-wetted states for either type of
heterogeneous inclusion shows that variation in wetting phase significantly affects the
form of the $S_{bw}$ dependence. It should be noted that in both cases the water-wetted matrix
parameters were used; it is expected that rock matrix parameters should vary with
changing wetting conditions. However, this effect cannot be quantified at present.

Let us now consider the relationship between pore fluid configuration and the elastic
behavior of the model. In the derivation of the mathematical formulations, it was shown
that the contents of a heterogeneous inclusion can be replaced by an effective pore fluid
having elastic moduli that are independent of the geometrical configuration of the pore
fluids within the inclusion. Therefore, it is only necessary to compare models containing
homogeneous and these "generic" heterogeneous inclusions when analyzing the effect of
pore fluid configuration on the macroscopic elastic behavior. Figures (4.11)-(4.14)
illustrate the computed values of $k^*$, $\mu^*$, $V_p^*$ and $V_s^*$, respectively, as a function of $S_{bw}$ for
models containing homogeneous inclusions. Models employing homogeneous inclusions
exhibit nearly linear variations of elastic properties with changing $S_{bw}$. Figures (4.15)-
(4.18) show the computed values of $k^*$, $\mu^*$, $V_p^*$ and $V_s^*$, respectively, as a function of $S_{bw}$
for models containing the "generic" heterogeneous inclusions. These results differ
significantly from those obtained with homogeneous inclusions in that distinctly non-
linear relationships are obtained between the elastic properties and $S_{bw}$. In particular, little change occurs in the elastic properties until high levels of $S_{bw}$ (i.e. in excess of 0.95) are obtained; this is when a significant increase occurs in the bulk modulus of the effective pore fluid.

Let us now consider the effects of local variation in saturation conditions on the effective physical properties by using three simple schemes for varying local $S_{bw}$ within the pore geometry spectrum. These simple schemes for increasing overall $S_{bw}$ from 0.00 to 1.00 are as follows: (1) water preferentially fills the largest $\alpha$ geometries; (2) water preferentially fills the smallest $\alpha$ geometries; and, (3) all geometries fill in unison. In all cases, homogeneous inclusions will be used to represent the individual pore spaces. These schemes are schematically shown in Figure (4.19). The pore spectrum used is an estimate for the pore geometries present within a Berea sandstone based upon Cheng (1978) and is given in Figure (4.20).

Figures (4.21)-(4.26) show the computed $\kappa^*_a$, $\sigma^*_a$, $k^*$, $\mu^*$, $V^*_p$ and $V^*_s$, respectively, as a function of $S_{bw}$ when these three schemes are used to vary the pore-scale fluid distribution. These results illustrate the significant differences which can occur in the functional relation between a given physical property and $S_{bw}$ when the sequence in which local saturation conditions vary is changed. In particular, changes in the saturation conditions in the smaller aspect ratio geometries have an effect on the dielectric and elastic behavior which is well out of proportion to their percentage of the total porosity (i.e. at high $S_{bw}$ value for scheme 1 and at low $S_{bw}$ values for scheme 2).

Unlike the other effective physical properties in this study, both $V^*_p$ and $V^*_s$ are
themselves functions of effective quantities. Therefore, variations in both \( V^*_p \) and \( V^*_s \) can be viewed in terms of the relative changes in these quantities. From Equations (4.11) and (4.12), it can be shown that

\[
\frac{dV^*_p}{dS_{bw}} = \frac{dV^*_p}{dk^*} \frac{dk^*}{dS_{bw}} + \frac{dV^*_p}{d\mu^*} \frac{d\mu^*}{dS_{bw}} + \frac{dV^*_p}{dp^*} \frac{dp^*}{dS_{bw}}
\]

\[
= \frac{1}{\rho^*} \left[ \frac{1}{2V^*_p} \frac{dk^*}{dS_{bw}} + 2 \frac{d\mu^*}{dS_{bw}} - \frac{dV^*_p}{V^*_p} \frac{\phi_{pore}}{2} \left( \rho^{(w)} - \rho^{(a)} \right) \right]
\]

(4.28)

and

\[
\frac{dV^*_s}{dS_{bw}} = \frac{dV^*_s}{d\mu^*} \frac{d\mu^*}{dS_{bw}} + \frac{dV^*_s}{dp^*} \frac{dp^*}{dS_{bw}}
\]

\[
= \frac{1}{\rho^*} \left[ \frac{1}{2V^*_s} \frac{d\mu^*}{dS_{bw}} - \frac{\phi_{pore}}{2} \frac{V^*_s}{2} \left( \rho^{(w)} - \rho^{(a)} \right) \right],
\]

(4.29)

where \( \rho^{(w)} \) and \( \rho^{(a)} \) are the densities of water and air, respectively. In general,

\[
\frac{dk^*}{dS_{bw}} \geq 0
\]

(4.30)

and

\[
\frac{d\mu^*}{dS_{bw}} \geq 0.
\]

(4.31)

Therefore, it follows that
\[
\frac{1}{2V_p^*} \frac{dk^*}{dS_{bw}} + \frac{2}{3V_p^*} \frac{d\mu^*}{dS_{bw}} \geq 0
\] 

(4.32)

and

\[
\frac{1}{2V_s^*} \frac{d\mu^*}{dS_{bw}} \geq 0 .
\] 

(4.33)

Thus, the existence of a relative minimum or maximum indicates values of \(S_{bw}\) where there is a change in the nature of the dominant mechanism that controls the relationship between \(V_p^*\) or \(V_s^*\) and \(S_{bw}\). A positive slope indicates that changes in the elastic moduli dominate; a negative slope means that variation in the effective density is controlling parameter. In general, variations in elastic wave velocities due changes in the saturation conditions in the larger \(\alpha\) pores are controlled by changes in \(\rho^*\). Conversely, changes in the saturation conditions in the smaller \(\alpha\) inclusions cause variations in \(V_p^*\) and \(V_s^*\) that are dominated by changes in the effective elastic moduli.

Conclusions

Inclusion-based formulations describing the elastic and electromagnetic wave propagation in a partially saturated porous rock have been derived by depicting this system as a rock matrix containing inclusions representing the individual pore spaces. By considering various styles of pore-scale fluid distribution, it has been found that drastically different functional relationships between these physical properties and overall saturation level are obtained while maintaining fixed constituent properties and pore
structure. Numerical simulations performed with these formulations have illustrated the significance of pore-scale fluid distribution in determining elastic and electromagnetic wave propagation within partially saturated porous rocks. In particular, it has been demonstrated that the pore fluid configuration within the individual pore spaces and the pore geometries in which local saturation conditions are varying are critical factors in determining the physical properties of partially saturated rocks. Further, changes in the saturation conditions in the smaller aspect ratio geometries have an affect on the physical response of the porous medium that is well out of proportion to their percentage of the total porosity. In the next chapter, these basic results will be used to analyze the experimentally measured dielectric and elastic response of partially saturated porous rocks.
<table>
<thead>
<tr>
<th></th>
<th>Rock Matrix</th>
<th>Deionized Water</th>
<th>Air</th>
</tr>
</thead>
<tbody>
<tr>
<td>$\kappa_a$</td>
<td>7.604</td>
<td>80.</td>
<td>1.01</td>
</tr>
<tr>
<td>$\sigma_a$ (in $\frac{S}{m}$)</td>
<td>$1.387 \times 10^{-3}$</td>
<td>$7.75 \times 10^{-4}$</td>
<td>0.</td>
</tr>
</tbody>
</table>

**Table 4.1:** Dielectric constants and conductivities used for the constituents of the model sandstone.

<table>
<thead>
<tr>
<th></th>
<th>Rock Matrix</th>
<th>Deionized Water</th>
<th>Air</th>
</tr>
</thead>
<tbody>
<tr>
<td>$k$ (in $\frac{N}{m^2}$)</td>
<td>$3.2 \times 10^{10}$</td>
<td>$2.237 \times 10^{9}$</td>
<td>$1.550 \times 10^{5}$</td>
</tr>
<tr>
<td>$\mu$ (in $\frac{N}{m^2}$)</td>
<td>$2.45 \times 10^{10}$</td>
<td>0.</td>
<td>0.</td>
</tr>
<tr>
<td>$\rho$ (in $\frac{kg}{m^3}$)</td>
<td>2650.</td>
<td>998.</td>
<td>1.293</td>
</tr>
</tbody>
</table>

**Table 4.2:** Elastic moduli and densities used for the constituents of the model sandstone.
Figure 4.1: Computed $\kappa_a^*/\kappa_z^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing homogeneous pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.2: Computed $\sigma^*_n/\sigma_n(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing homogeneous pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.3: Computed $\kappa_a^* / \kappa_a^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing pores filled with the water-wet effective pore fluid. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.4: Computed $\frac{\sigma_a^*}{\sigma_b^*}(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing pores filled with the water-wet effective pore fluid. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.5: Computed $\kappa_n^* / \kappa_a^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing water-wet confocally-layered pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.6: Computed $\sigma_n^*/\sigma_n^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing water-wet confocally-layered pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.7: Computed $\kappa_a^*/\kappa_a^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing pores filled with the air-wet effective pore fluid. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.8: Computed $\kappa^*/\kappa^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing air-wet confocally-layered pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.9: Computed $\sigma_a^*/\sigma_a^*(S_{BW}=0.00)$ as a function of $S_{BW}$ for a series of simple models containing pores filled with the air-wet effective pore fluid. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.10: Computed $\sigma_a^*/\sigma_a(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing air-wet confocally-layered pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.11: Computed $k^*/k^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing homogeneous pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.12: Computed $\mu^*/\mu^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing homogeneous pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.13: Computed $V_p^*/V_p^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing homogeneous pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.14: Computed $V_s^*/V_s^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing homogeneous pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.15: Computed $k^*/k^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing heterogeneous pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.16: Computed $\mu^*/\mu^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing heterogeneous pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.17: Computed $V_p^*/V_p(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing heterogeneous pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.18: Computed $V_s^*/V_s^*(S_{bw}=0.00)$ as a function of $S_{bw}$ for a series of simple models containing heterogeneous pores. Aspect ratio ($\alpha$) of the pore geometry present in a given model is denoted.
Figure 4.19: Schematic diagram of the three fluid distribution schemes using homogeneous pores for varying local $S_{bw}$ level within the pore geometry spectrum. Pore space is illustrated with three different $\alpha$ spheroids. Shading=water; white=air. Scheme one-water preferentially fills large $\alpha$ pores; scheme two-water preferentially fills small $\alpha$ pores; scheme three-water fills all pore geometries in unison.
Figure 4.20: Pore geometry spectrum in volume fraction for model sandstone.
**Figure 4.21:** Computed $\kappa_\alpha^*$ as a function of $S_{bw}$ for the model sandstone using the three fluid distribution schemes for varying local $S_{bw}$ level.
Figure 4.22: Computed $\sigma_\Phi^*$ as a function of $S_{bw}$ for the model sandstone using the three fluid distribution schemes for varying local $S_{bw}$ level.
Figure 4.23: Computed $k^*$ as a function of $S_{bw}$ for the model sandstone using the three fluid distribution schemes for varying local $S_{bw}$ level.
Figure 4.24: Computed $\mu^*$ as a function of $S_{bw}$ for the model sandstone using the three fluid distribution schemes for varying local $S_{bw}$ level.
Figure 4.25: Computed $V_p^*$ as a function of $S_{bw}$ for the model sandstone using the three fluid distribution schemes for varying local $S_{bw}$ level.
Figure 4.26: Computed $V_s^*$ as a function of $S_{bw}$ for the model sandstone using the three fluid distribution schemes for varying local $S_{bw}$ level.
Chapter Five

The Effects of Pore-Scale Fluid Distribution on Elastic and Electromagnetic Wave Propagation in Partially Saturated Porous Rocks: An Analysis of Experimental Results

Introduction

A number of experimental studies have clearly demonstrated the dependence of macroscopic physical properties on pore-scale fluid distribution (Domenico, 1976, 1977; Bourbie and Zinszner, 1984; Knight and Nur, 1987a; Longeron et al., 1989; Knight and Nolen-Hoeksema, 1990); differing styles of pore-scale fluid distribution were obtained in these laboratory experiments by changing the saturation technique. The theoretical importance of pore-scale fluid distribution in determining the dielectric and elastic response of a partially saturated porous rocks was demonstrated in the last chapter. The mathematical formulation employed in that analysis will now be used to study the experimentally measured elastic wave velocities and apparent dielectric constant of partially saturated porous materials.

Of particular interest are the results of Knight and Nur (1987a) and Knight and Nolen-Hoeksema (1990) that give the measured apparent dielectric constant and elastic wave velocities, respectively, for similar samples of a tight gas sandstone undergoing the same two saturation techniques. Before the existence of these results, individual experimental studies have been concerned with the effect of differing saturation techniques on a single
macroscopic physical property. Further, a variety of different saturation procedures and porous media samples were used in these studies; hence, there is difficulty in comparing the results of these experiments. The data set obtained by Knight and co-workers permits a direct assessment of the variation in the dielectric and elastic response of the same rock type undergoing the same saturation methods.

Further analysis of the relationship between elastic wave velocities and differing styles of pore-scale fluid distribution is undertaken by considering the experimental data obtained by Domenico (1977). In his experiment, Domenico subjected a glass bead packing to two saturation techniques that differ from those employed in the tight gas sandstone experiments. It is shown in this study that the physical properties of both the tight gas sandstone and the glass bead packing are replicated by utilizing simple models to describe the geometrical configuration of the constituents. In addition, it is seen that a single geometrical model for pore-scale fluid distribution can be used simultaneously to estimate the dielectric and elastic behavior of the tight gas sandstone.

Tight Gas Sandstone: Experimental Results

Knight and Nur (1987a) measured the apparent dielectric constant \( \kappa_a \) of a tight gas sandstone sample from the Spirit River Formation in the Alberta Basin; Knight and Nolen-Hoeksema (1990) measured the compressional wave velocity \( V_p \) and shear wave velocity \( V_s \) on a similar sample from the same formation. The experimental procedures employed are described in detail in these two studies. The sample used in the dielectric measurements has a porosity of 0.070 and a gas permeability of 7.18 \( \mu \) D; the sample used in the study of the elastic wave velocities has a porosity of 0.052 and a permeability of 1 \( \mu \) D. A complete
petrographic description of the Spirit River samples is found in Walls (1982).

In both laboratory studies, measurements were made during a continuous imbibition-drainage experiment (where imbibition is the displacement of air by water in a water-wet system and drainage is the displacement of water by air). Starting with a desiccator dried sample, $S_w$ was initially increased through imbibition by adsorption of water vapor, followed by the soaking of the sample in deionized water. The maximum value of $S_w$ obtained through this procedure was approximately $S_w = 0.90$ in both experiments. Water saturation was subsequently decreased through drainage by evaporative drying. As $S_w$ was varied, its value was determined by means of sample weight.

Figures (5.1)-(5.3) display the measurements of $\kappa_a$, $V_p$ and $V_s$, respectively, obtained during the imbibition-drainage cycle. The dielectric data were measured at a frequency of 105 kHz; the central frequencies for the $V_p$ and $V_s$ measurements were 1 MHz and 600 kHz, respectively. Each data set clearly illustrates the dependence of the functional relationship between a physical property and $S_w$ on the saturation technique used. There are clearly two distinct regions in both the apparent dielectric constant and elastic wave velocity data sets that can be related to the location and nature of the water present in the pore space. Knight and Nur (1987a,b) defined a critical value of $S_w$ ($S_w^0$) separating these two regions in the dielectric response (i.e. $S_w^0 = 0.15$); these two regions are labelled in Figure (5.1). In region 1 ($S_w < S_w^0$), the water present in the rock exists as monolayers of adsorbed water coating the surfaces of the pore space. In region 2 ($S_w > S_w^0$), a bulk water phase filling the central portion of the pore space is also present. For the sample used in the dielectric measurements, $S_w^0$ has been shown to correspond to a surface layer.
of bound water, approximately 1 nm in thickness (Knight and Nur, 1987a,b). The same
distinction can be made in an analysis of the elastic data shown in Figures (5.2) and (5.3).
In this case, the value of \( S_w^0 \) cannot be precisely defined and a zone of transition in which
\( S_w^0 \) occurs is denoted (i.e. \( 0.25 \leq S_w^0 \leq 0.35 \)).

In region 1, where only bound water is present, a significant increase in \( \kappa_a \) and a
decrease in the elastic wave velocities, particularly \( V_s \) (Clark et al., 1980), occurs as \( S_w \)
increases towards \( S_w^0 \). Variations in the physical properties within this saturation range are
due to complex rock-bound water interactions. Since rocks in situ usually contain the
bound water phase, the physical behavior of rocks in this region of \( S_w \) values will not be
considered in this study. Of interest in this study is the theoretical analysis of the hysteresis
that occurs in both data sets at higher \( S_w \) values. This hysteresis in the physical properties
has been attributed to changes in the geometrical distribution of fluids within the pore space
that occur during the imbibition–drainage cycle (Knight and Nur, 1987b; Knight and
Nolen-Hoeksema, 1990). Based on earlier works by Haines (1930) and Foster (1932), the
following model was proposed by Knight and Nur (1987a) in order to explain the results
of their dielectric measurements.

At low levels of \( S_w \), as water content is increased by the adsorption of water vapor, the
water preferentially adsorbs on the surfaces of the pore space. This adsorption leads to the
development of thick surface layers of water along the outer portion of the pore volume
with a thin, interconnected gas phase within the central portion. This arrangement of gas
and water means that the individual pore spaces are in a partially saturated state. When the
imbibition process is continued at higher levels of \( S_w \) by soaking the sample, the bulk water
phase tends to maintain this established geometrical configuration by continuing to move
along and coat the surfaces. However, this geometrical configuration is a metastable state persisting until closure occurs between the surface layers of bulk water. At this point, the rearrangement of the bulk water and air to a more thermodynamically stable configuration occurs, causing the break-up of the interconnected gas phase and the filling of the central pore volumes with the bulk water phase.

The model proposed for the drainage process is quite different. During drainage, the bulk water empties from each individual pore space as it is accessed, removing all but the first few tightly-held monolayers of bound water. Hence, the geometrical configuration of the pore fluids that occurs during imbibition is not re-created during drainage. The imbibition process favors the development of pores with co-existing gas and bulk water phases; the drainage process tends to develop either fully bulk water-saturated or fully gas-saturated pore spaces.

Since the object of this study is to assess the effects associated with the geometrical distribution of the bulk water phase, we are only concerned with the behavior of the tight gas sandstone in the range of the $S_w$ values between $S_w^0$ and 1.00. In this region, the discussion will be in terms of bulk water saturation ($S_{bw}$) over a range from 0.00 to 1.00. Further, the bound water phase is considered to be part of a wetted rock matrix system (Knight and Endres, 1990); this implies that the total porosity of the samples must also be corrected for the existence of the bound water phase. If the value of $S_w^0$ is taken to be 0.15 and 0.35 for the dielectric and the elastic measurements, respectively, then the corrected porosity ($\phi_c$, the total porosity occupied by the bulk water phase) for the dielectric and elastic samples is 0.0595 and 0.0338, respectively.
Tight Gas Sandstone: Numerical Simulation

Let us first consider the drainage process. Given that the segregation of pore fluids occurs, homogeneous inclusions composed of either air or bulk water are used to model the pore-scale fluid distribution. The effective permittivity and elastic moduli of such a system with homogeneous pore spaces is calculated by using Equation (4.4) and Equations (4.7) and (4.8), respectively. The overall value of $S_{bw}$ is varied by simultaneously changing the local $S_{bw}$ level in all pore geometries such that emptying occurs in unison across the entire pore geometry spectrum. This scenario for the pore-scale fluid distribution is schematically shown in Figure (5.4). Given that the topology of the pore space within a rock necessitates essentially simultaneous drainage of all aspect ratio geometries in order to access the total pore space, this scheme is considered to be realistic for modeling the experimental data.

For the imbibition process, it is necessary to treat each individual pore as being simultaneously occupied by air and bulk water until the rearrangement of pore fluids occurs. Given the configuration of thick surface layers of bulk water and an interconnected air phase in the central volume of the pores, confocally-layered spheroids consisting of a bulk water outer shell and an inner air spheroid are reasonable inclusion geometries to use in simulating the process. Further, it will be assumed that all pores with a given shape will have identical levels of local $S_{bw}$ prior to the critical rearrangement phenomenon. The effective permittivity for a system containing confocally-layered pore spaces is computed from Equation (4.21). Since the pore spaces are heterogeneous, Equations (4.16) through (4.18) are used to determine the effective elastic moduli.

To simulate the pore-scale fluid distribution which occurs during the imbibition process prior to the critical rearrangement of pore fluids, the following idealized geometrical
arrangement of pore fluids was employed; this geometrical configuration is schematically illustrated in Figure (5.5). As $S_{bw}$ starts to increase from 0.00, the lower aspect ratio geometries, which are generally associated with the pore throats (in the case of the tight sandstone model, this will be assumed to be $\alpha \leq 0.5$), are initially filled to some level of local $S_{bw}$. The local levels of $S_{bw}$ specified for the various pore throat shapes in this simulation are given in Table (5.1). This establishes the presence of the narrow gas pockets within the pore throats. This geometrical configuration within the pore throats is maintained as the overall level of $S_{bw}$ of the medium is increased by confocally filling the remaining pore geometries until $S_{bw}^c$, the point at which the rearrangement of pore fluids occurs, is achieved. The critical value of overall water saturation $S_{bw}^c$ is taken to be 0.65 and 0.80 in the case of the dielectric and velocity models, respectively. After the rearrangement of the pore fluids, it will be assumed that the pore-scale fluid distribution is identical to that used to model the drainage process at the given value of $S_{bw}$.

The constituent properties used for modeling the dielectric and elastic properties of the tight gas sandstone are given in Tables (5.2) and (5.3), respectively. The frequency used in the computation of the dielectric properties is 105 kHz. Figure (5.6) illustrates the pore geometry spectra used to describe the two samples. These spectra are based on the work of Toksöz et al. (1976) and Cheng and Toksöz (1979). Since different samples were used in each of the experiments, small differences in the fractional volume for a given aspect ratio are permitted between the pore geometry spectra representing the two samples. In both cases, the total porosity used is corrected for the bound water being incorporated as part of the rock matrix system.
Figure (5.7) shows the computed apparent dielectric constant ($\kappa^*_a$); Figures (5.8) and (5.9) show the computed compressional and shear wave velocities ($V^*_p$ and $V^*_s$), respectively. In all cases, the experimental data is superimposed on the model response. While there are differences between the theoretical models and experimental data, the basic form of both are in good agreement. The model accurately predicts the functional form and magnitude of the hysteresis which occurs in both the apparent dielectric constant and elastic wave velocities as the sample undergoes the imbibition - drainage cycle. The differences which are observed can be attributed to inaccuracies both in the geometrical schemes used to represent the pore-scale fluid distribution and in the parameters used for the constituents and pore geometry spectra. However, it is significant that the relatively simple scenarios used to depict the geometrical configurations occurring within the rock give good results. In particular, it can be observed that the existence of the thin interconnected air phase throughout the pore space during the imbibition technique prior to the rearrangement of the pore fluids causes a significant enhancement in $\kappa_a$ and a depressed, almost constant value of $V_p$ and $V_s$. Further, the drainage technique gives a much more uniform variation in physical properties with changes in $S_{bw}$.

**Glass Bead Packing: Experimental Data**

Domenico (1976, 1977) investigated the functional relationship between elastic wave velocities and $S_w$ for a glass bead packing by employing two different saturation methods that he referred to as the flow and imbibition techniques. The experimental procedure used is given in detail in the above references and is briefly summarized below. In the flow technique, the sample was initially saturated with nitrogen gas; the gas was subsequently
totally displaced by a brine solution. Partial saturation was achieved by injecting a series of
gas-brine mixtures into the sample. The gas and brine were mixed prior to injection;
however, the two pore fluids remained immiscible during the emplacement procedure.
Approximately ten pore volumes of the prepared fluid mixture were passed through the
glass bead packing between measurements.

The imbibition method initially utilized the flow method to obtain a partially saturated
state. Subsequently, the fluid exit valve was closed and pure brine was injected into the
glass bead packing until the gas phase present within the pore space was completely
dissolved into solution. This results in a significantly increased pore fluid pressure ($p_f$)
within the porous medium. To obtain a partially saturated medium, $p_f$ was decreased and
gas was exsolved from solution. During this process, the confining pressure ($p_c$) was
adjusted such that the effective pressure ($p_e = p_c - p_f$) was constant for all measurements.
During both saturation techniques, the level and uniformity of $S_w$ within the glass bead
packing was monitored by means of an X-ray system.

Figures (5.10) - (5.13) illustrate the experimental results for the elastic wave velocities
as a function of $S_w$; these measurements were performed at a central frequency of 200
kHz. Over the range of $S_w$ values between 0.70 and 1.00, the two saturation techniques
gave distinctly different results for $V_p$ at both effective pressures. Although the flow
technique produced a less uniform distribution of $S_w$ values along the length of the sample,
these variations are not sufficient to explain the observed differences between the two
saturation methods. For $S_w$ values between 0.88 and 1.00, the flow technique gives a very
uniform $S_w$ distribution while differences between the two methods are the most
pronounced. There are indications in the $V_s$ measurements that the two saturation
techniques are different; however, the data are not conclusive.

It is hypothesized in this study that Domenico's flow technique results in a pore-scale
fluid distribution similar to that obtained during the imbibition method employed by Knight
and co-workers. By initially saturating the medium completely with brine, it is established
as the wetting phase throughout the glass bead packing. Subsequent injection of a distinct
gas phase results in it being placed in the central portion of the pore spaces. Up to some
critical level of $S_w$ ($S_w^c$), the pressure within the gas can overcome other thermodynamic
factors in order to establish and maintain an interconnected gas phase. Once again, such a
configuration of fluids is thermodynamically metastable; and, gas pressure cannot maintain
gas interconnectivity beyond $S_w^c$.

However, there is a significant difference between these two saturation techniques.
Knight's method results in a continuous sequence of changes in the pore-scale distribution
of fluids which represents a single realization of the metastable process. Hence, physical
properties display a clear functional relationship as $S_w$ is varied. However, this metastable
process is sensitive to the saturation history of the porous medium. By flushing large
volumes of fluids through the porous medium, Domenico's technique produces relatively
independent realizations of the metastable process for each measurement. Hence, this
technique can produce significantly different pore-scale fluid distributions for similar values
of $S_w$. In particular, the value of $S_w$ at which $S_w^c$ occurs varies between realizations of this
process. This is reflected in the scatter observed in the elastic wave velocity data and the
variation observed in local $S_w$ (see Figure 22 in Domenico, 1977) over the range of $S_w$
values between 0.70 and 0.90.
Domenico's imbibition technique results in a style of pore-scale fluid distribution which is significantly different from those previously discussed. The exsolution of a gas phase from solution to obtain high values of $S_w$ should result in the formation of individual gas bubbles throughout the pore space. Further, the creation of the gas bubbles should be relatively uniform through the pore structure of the rock.

**Glass Bead Packing: Numerical Simulation**

The pore-scale fluid distribution resulting from Domenico's flow method is simulated by means of the same geometrical scheme used to represent Knight's imbibition technique. The generation of gas bubbles implies that Domenico's imbibition technique is characterized by heterogeneous pore spaces. Hence, Equations (4.16) through (4.18) describe the effective elastic moduli resulting from this saturation technique. Further, it will be assumed that the exsolution process produces a uniform distribution of gas bubbles such that the local $S_w$ within a given pore space is the same as the overall value of $S_w$.

The elastic properties of the constituents are given in Table (5.4); Figure (5.14) gives the pore geometry spectrum used to represent the glass bead packing. It is assumed that the relative amount of the bound water phase is small in comparison to that of the bulk water phase within the glass bead packing. Therefore, corrections for the volume occupied by the bound water phase were not made; and, the results of the numerical simulation are expressed in terms of total water saturation. In simulating the flow technique, $S_{wc}$ was taken to be 0.80; and, pore throat geometries were assumed to have aspect ratios of 0.5 or less.
Figure (5.15) and (5.16) show the computed compressional and shear wave velocities (\(V_p^*\) and \(V_s^*\)), respectively, for the glass bead packing model. The functional relationships between \(V_p\) and \(S_w\) obtained from the numerical simulation for the two saturation techniques is in good qualitative agreement with the experimental data. In particular, the dominance of the density effect over the lower \(S_w\) range during the flow method and the dramatic decrease in \(V_p\) as \(S_w\) decreases from full water saturation during the imbibition technique are accurately reproduced. Numerical results for \(V_s\) show that the expected difference between the two saturation methods is very small. Hence, it is not surprising that the experimental data is not conclusive.

Conclusions

The processes involved in the filling and emptying of a porous rock with fluids are undoubtedly complex. In this study, simple models are employed to describe the pore-scale fluid distribution resulting from the use of a given saturation technique. The treatment of these saturation methods has been simplified to consider only the basic geometrical elements characterizing the pore space and the contained fluids. The good agreement that is obtained between the theoretical models and the experimental data implies that these geometrical effects are critical factors in determining the physical properties of partially saturated rocks. Further, it is significant that a single geometrical model for the pore-scale fluid distribution was used simultaneously to estimate the elastic wave velocities and apparent dielectric constant of a partially saturated rock.

In the theoretical formulations used, it was assumed that communication between individual pore spaces is minimal. Hence, intrapore phenomena control both the elastic and
electromagnetic wave propagation in the porous medium. This is achieved experimentally by taking measurements at sufficiently high sampling frequencies and/or using porous media that possess adequately low permeability. In the cases of the tight gas sandstone and the glass bead packing, these requirements appear to have been satisfied.

However, a review of the experimental work in the literature on elastic wave velocities in partially saturated rocks suggests that these conditions commonly do not hold. Drainage by evaporative drying has been used to control overall $S_w$ in a number of studies (Wyllie et al., 1956; Gregory, 1976; Murphy, 1982b, 1984, 1985). The theoretical results for the tight gas sandstone (Figures 5.8 and 5.9) and Berea sandstone (Figure 4.25 and 4.26) imply that elastic wave velocities should vary in a nearly linear manner with changing $S_w$. Results obtain by Murphy (1982b, 1984, 1985) show that significant departure from this quasi-linear relationship occurs as sampling frequency decreases. In the case of Massilon sandstone (Murphy, 1982b), the combination of very high permeability (737 mD) and sampling frequencies in the vicinity of 500 Hz gives results that are well described by the phenomenological formulations of Gassmann (1951b) and Biot (1956) where complete pressure communication within the pore fluid is assumed. However, the behavior of tight gas sandstones (Murphy, 1984) and granite (Murphy, 1985) measured at sampling frequencies of approximately 5 kHz display an intermediate type of functional relationship that cannot be adequately described by the inclusion-based formulations used in this study or the phenomenological approach. Further, there appear to be no adequate methods to determine a priori whether a particular formulation is valid for a given porous rock for a specific set of conditions.
<table>
<thead>
<tr>
<th>Range of $\alpha$</th>
<th>Local Level of $S_{bw}$</th>
</tr>
</thead>
<tbody>
<tr>
<td>$5 \times 10^{-1} \geq \alpha &gt; 10^{-1}$</td>
<td>0.98</td>
</tr>
<tr>
<td>$10^{-1} &gt; \alpha \geq 10^{-2}$</td>
<td>0.75</td>
</tr>
<tr>
<td>$10^{-2} &gt; \alpha \geq 10^{-3}$</td>
<td>0.20</td>
</tr>
<tr>
<td>$10^{-3} &gt; \alpha \geq 10^{-4}$</td>
<td>0.10</td>
</tr>
</tbody>
</table>

**Table 5.1:** Level of local $S_{bw}$ maintained in the pore throat geometries during the imbibition process prior to the rearrangement phenomenon.

<table>
<thead>
<tr>
<th>Rock Matrix</th>
<th>Deionized Water</th>
<th>Air</th>
</tr>
</thead>
<tbody>
<tr>
<td>$\kappa_a$</td>
<td>13.75</td>
<td>80.</td>
</tr>
<tr>
<td>$\sigma_a$ (in $\text{S/m}$)</td>
<td>$2.53 \times 10^{-3}$</td>
<td>$5. \times 10^{-5}$</td>
</tr>
</tbody>
</table>

**Table 5.2:** Dielectric constants and conductivities used for the constituents of the tight gas sandstone model.
<table>
<thead>
<tr>
<th></th>
<th>Rock Matrix</th>
<th>Deionized Water</th>
<th>Air</th>
</tr>
</thead>
<tbody>
<tr>
<td>$k \left( \frac{N}{m^2} \right)$</td>
<td>$2.9 \times 10^{10}$</td>
<td>$2.237 \times 10^{9}$</td>
<td>$1.550 \times 10^{5}$</td>
</tr>
<tr>
<td>$\mu \left( \frac{N}{m^2} \right)$</td>
<td>$2.6 \times 10^{10}$</td>
<td>0.</td>
<td>0.</td>
</tr>
<tr>
<td>$\rho \left( \frac{kg}{m^3} \right)$</td>
<td>2650.</td>
<td>998.</td>
<td>1.293</td>
</tr>
</tbody>
</table>

**Table 5.3:** Elastic moduli and densities used for the constituents of the tight gas sandstone model.

<table>
<thead>
<tr>
<th></th>
<th>Glass Bead Packing</th>
<th>Brine</th>
<th>Gas</th>
</tr>
</thead>
<tbody>
<tr>
<td>$k \left( \frac{N}{m^2} \right)$</td>
<td>$2.137 \times 10^{10}$</td>
<td>$2.098 \times 10^{9}$</td>
<td>$1.378 \times 10^{6}$</td>
</tr>
<tr>
<td>$\mu \left( \frac{N}{m^2} \right)$</td>
<td>$3.053 \times 10^{9}$</td>
<td>0.</td>
<td>0.</td>
</tr>
<tr>
<td>$\rho \left( \frac{kg}{m^3} \right)$</td>
<td>2420.</td>
<td>1080.</td>
<td>11.67</td>
</tr>
</tbody>
</table>

**Table 5.4:** Elastic moduli and densities used for the constituents of the glass bead packing model.
Figure 5.1: Measured apparent dielectric constant ($\kappa_a$) of a tight gas sandstone undergoing imbibition-drainage cycle (sampling frequency=105kHz) (from Knight and Nur, 1987a). Region (1): only bound water phase present; region (2): bulk water present.
Figure 5.2: Measured compressional wave velocity ($V_p$) of a tight gas sandstone undergoing imbibition-drainage cycle (from Knight and Nolen-Hoeksema, 1990). Region (1): only bound water phase present; region (2): bulk water present.
Figure 5.3: Measured shear wave velocity ($V_s$) of a tight gas sandstone undergoing imbibition-drainage cycle (from Knight and Nolen-Hoeksema, 1990). Region (1): only bound water phase present; region (2): bulk water present.
Figure 5.4: Schematic diagram of the scenario used to simulate the pore-scale fluid distribution during the drainage process: all pore geometries emptying in unison. Black=water, white=air.
Figure 5.5: Schematic diagram of the scenario used to simulate the pore-scale fluid distribution during the imbibition process: (a) initial filling of the pore throats; (b) filling of remaining pore geometries to the critical overall $S_{bw}$ value; (c) all pore geometries filling in unison after rearrangement phenomena. Black=water, white=air.
Figure 5.6: Pore geometry spectra (in volume fraction) used for dielectric and elastic models of the tight gas sandstone.
Figure 5.7: Computed apparent dielectric constant ($\kappa_a^*$) as a function of bulk water saturation ($S_{bw}$) for the tight gas sandstone model using the imbibition-drainage scenarios. Experimental data are superimposed on the theoretical results.
Figure 5.8: Computed compressional wave velocity ($V_p^*$) as a function of bulk water saturation ($S_{bw}$) for the tight gas sandstone model using the imbibition-drainage scenarios. Experimental data are superimposed on the theoretical results.
Figure 5.9: Computed shear wave velocity ($V_s^*$) as a function of bulk water saturation ($S_{bw}$) for the tight gas sandstone model using the imbibition-drainage scenarios. Experimental data are superimposed on the theoretical results.
Figure 5.10: Measured compressional wave velocity ($V_p$) of a glass bead packing undergoing flow and imbibition methods ($p_e=10.33$ MPa) (from Domenico, 1977).
Figure 5.11: Measured shear wave velocity ($V_s$) of a glass bead packing undergoing flow and imbibition methods ($p_c=10.33$ MPa) (from Domenico, 1977).
Figure 5.12: Measured compressional wave velocity ($V_p$) of a glass bead packing undergoing flow and imbibition methods ($p_e=31.00$ MPa) (from Domenico, 1977).
Figure 5.13: Measured shear wave velocity ($V_s$) of a glass bead packing undergoing flow and imbibition methods ($p_c=31.00$ MPa) (from Domenico, 1977).
Figure 5.14: Pore geometry spectrum (in volume fraction) used for elastic models of the glass bead packing.
Figure 5.15: Computed compressional wave velocity ($V_p^*$) as a function of total water saturation ($S_w$) for the glass bead packing model using scenarios for Domenico's imbibition and flow methods.
Figure 5.16: Computed shear wave velocity ($V_s^*$) as a function of total water saturation ($S_w$) for the glass bead packing model using scenarios for Domenico's imbibition and flow methods.
Chapter Six

A Model for Incorporating the Effects of Surface Phenomena in the
Dielectric Response of Heterogeneous Media

Introduction

The presence of surface phenomena can significantly affect the dielectric response of a heterogeneous medium. One method for incorporating the effect of surface phenomena in a theoretical formulation is to redefine the bulk components. In particular, the adsorbed water phase present within a water-wetted sandstone was considered to be part of the rock matrix system; and, the parameters used to describe this system reflect this wetted state (Knight and Endres, 1990). However, this approach does not preserve the geometrical structure of the system; the effects of surface phenomena are uniformly distributed throughout the rock matrix. In order to retain the structure of the system, it is necessary to employ a mathematical formulation that localizes these mechanisms at the interfaces within the system.

The region in which surface phenomena occur is generally very thin in comparison to the dimensions of the individual pore spaces. Thus, the effects of surface phenomena can be incorporated into the analysis of an individual inclusion by specifying the appropriate boundary conditions. This approach was first used by O’Konski (1960) in order to introduce surface currents and an associated surface conduction into an analysis of a spherical colloidal particle. Subsequently, a number of studies (Schwarz, 1962; Schurr,
1964; Dukhin, 1971; Dukhin and Shilov, 1974; Fixman, 1980, 1983) have employed a variety of different boundary conditions in order to incorporate surface phenomena into the analysis of a spherical particle.

To adequately describe the geometry present within a porous medium, it is necessary to consider more versatile inclusion shapes, such as ellipsoids. However, analytical solutions for ellipsoidal inclusion geometries employing the necessary boundary conditions have not been found. O’Konski (1960) offered an approximate solution for the ellipsoidal case based on his results for a spherical particle. He assumed that the effects of the surface phenomena could be represented by an equivalent volumetric mechanism which is uniformly distributed throughout the ellipsoid. Again, this approach radically alters the geometrical configuration of the system under consideration.

In this chapter, a new approximation for the case of an ellipsoidal inclusion influenced by surface phenomena is present. Motivated by the technique used by Miles and Robertson (1932) in an analysis of the spherical case, the starting point of this model is a confocally-layered ellipsoidal system. The outer shell in which the surface phenomena take place is taken to be infinitesimally thin while the number of unit charge carriers and unit dipoles within this region is conserved. In the limit, this results in a homogeneous (i.e. single component) inclusion possessing a surface permittivity representing the surface phenomena. For spherical geometries, this method gives a result identical to that obtained rigorously by means of boundary conditions.

Response of a Homogeneous Inclusion with Surface Phenomena

Determination of the macroscopic dielectric properties of a heterogeneous medium by
means of an inclusion-based formulation requires an analysis of the response of the individual inclusions. The analysis undertaken to this point has employed boundary conditions that exclude surface phenomena at the interfaces between the inclusions and background material. It is possible to incorporate the effects of surface phenomena into this analysis by specifying the appropriate boundary conditions; however, an analytical solution has not been determined for the case of a general ellipsoidal inclusion.

An approximation for a homogeneous ellipsoidal inclusion when surface phenomena occur can be obtain by initially representing this system as a two-component confocally-layered ellipsoidal inclusion. The outer and inner ellipsoids of this confocal system have principal semi-axis lengths of \((a_1^{(1)}, a_2^{(1)}, a_3^{(1)})\) and \((a_1^{(2)}, a_2^{(2)}, a_3^{(2)})\), respectively. The relationship between the semi-axis lengths of the two ellipsoids is given by

\[
(a_j^{(1)})^2 = (a_j^{(2)})^2 + \xi \quad (j=1, 2, 3) \tag{6.1}
\]

where \(\xi\) is the ellipsoidal coordinate associated with confocal ellipsoids. The shape of the outer and inner ellipsoids are uniquely determined by the aspect ratio pairs \((\alpha_1^{(1)}, \alpha_2^{(1)})\) and \((\alpha_1^{(2)}, \alpha_2^{(2)})\), respectively. For a heterogeneous system containing this type of inhomogeneity, Equation (3.6) can be rewritten as

\[
\mathcal{V} \varepsilon^*_g \hat{E} = \mathcal{V} \varepsilon_g^{(0)} \hat{E} + \int_{V_{inc}} \left( \varepsilon_g(x) - \varepsilon_g^{(0)} \right) E^{(i)}(x) \, dv \tag{6.2}
\]

where the perturbation term now involves the volume integration over the confocally-
layered inclusions. Since the outer shell and inner ellipsoid are composed of homogeneous materials with generalized permittivities \( \varepsilon^{(1)}_g \) and \( \varepsilon^{(2)}_g \), respectively, the evaluation of the perturbation term only requires the determination of the electric field strength within each component.

As was done in Chapter Three, \( E^{(i)}(x) \) within an individual inclusion is estimated by treating it as an isolated inclusion embedded in an infinite homogeneous background with generalized permittivity \( \varepsilon^{(b)}_g \) subjected to a uniform applied electric field with strength \( E_a \). Stepin (1965) determined the electric field strength within such a system; his analysis showed that the electric field strength within each layer of the inclusion can be decomposed into "homogeneous" and "heterogeneous" components. Further, it was demonstrated that integration of the contribution due to the "heterogeneous" component over the volume of a given layer is zero. Hence, only the "homogeneous" component need be considered in the evaluation of the perturbation term. The "homogeneous" component of the electric field strength is

\[
E^{(1)}_{\text{homo}} = E_a \sum_{j=1}^{3} \left[ \frac{-\eta_{1j}(\varepsilon^{(1)}_g, \varepsilon^{(2)}_g, a(2))}{\eta_{0j}(\varepsilon^{(b)}_g, \varepsilon^{(1)}_g, \varepsilon^{(2)}_g, a(1), a(2))} \cos \theta_j \hat{x}_j \right] \tag{6.3}
\]

within the outer ellipsoidal shell and

\[
E^{(2)}_{\text{homo}} = E_a \sum_{j=1}^{3} \left[ \frac{1}{\eta_{0j}(\varepsilon^{(b)}_g, \varepsilon^{(1)}_g, \varepsilon^{(2)}_g, a(1), a(2))} \cos \theta_j \hat{x}_j \right] \tag{6.4}
\]
within the inner ellipsoid. The terms \( \eta_{0j} \) and \( \eta_{1j} \) are defined as follows:

\[
\eta_{0j}(\varepsilon_g^{(b)}, \varepsilon_g^{(1)}, \varepsilon_g^{(2)}, a^{(1)}, a^{(2)}) = \left[ A_j^{(2)}(\alpha_1^{(2)}, \alpha_2^{(2)}) - 1 \right] \frac{\varepsilon_g^{(2)}}{\varepsilon_g^{(1)}} A_j^{(2)}(\alpha_1^{(2)}, \alpha_2^{(2)}) \times
\left[ A_j^{(1)}(\alpha_1^{(1)}, \alpha_2^{(1)}) - 1 \right] - \frac{\varepsilon_g^{(1)}}{\varepsilon_g^{(b)}} A_j^{(1)}(\alpha_1^{(1)}, \alpha_2^{(1)}) + \frac{a_2^{(2)} a_3^{(2)}}{a_1^{(1)} a_2^{(1)} a_3^{(1)}} A_j^{(1)}(\alpha_1^{(1)}, \alpha_2^{(1)}) \\
\left[ 1 - A_j^{(1)}(\alpha_1^{(1)}, \alpha_2^{(1)}) \right] \left( 1 - \frac{\varepsilon_g^{(1)}}{\varepsilon_g^{(b)}} \right) \left( 1 - \frac{\varepsilon_g^{(2)}}{\varepsilon_g^{(1)}} \right) \right]
\tag{6.5}
\]

and

\[
\eta_{1j}(\varepsilon_g^{(1)}, \varepsilon_g^{(2)}, a^{(2)}) = \left[ A_j^{(2)}(\alpha_1^{(2)}, \alpha_2^{(2)}) - 1 \right] \frac{\varepsilon_g^{(2)}}{\varepsilon_g^{(1)}} A_j^{(2)}(\alpha_1^{(2)}, \alpha_2^{(2)}) , \tag{6.6}
\]

where \( A_j^{(1)}(\alpha_1^{(1)}, \alpha_2^{(1)}) \) and \( A_j^{(2)}(\alpha_1^{(2)}, \alpha_2^{(2)}) \) are the depolarization coefficients with respect to the \( j \)th axis for the outer and inner ellipsoid, respectively.

Using Equations (6.3) and (6.4), it can be shown that the expected value of the volume integral for a confocally-layered inclusion that is oriented in a uniformly random manner is given by

\[
\left\langle \int_{V_{inc}} (\varepsilon_g(x) - \varepsilon_g^{(0)}) E^{(i)}(x) dv \right\rangle = \frac{1}{3} \varepsilon_s e_g^{(b)} (a_1^{(1)} a_2^{(1)} a_3^{(1)}) \times
\]
\[
\begin{align*}
\left\{ \sum_{j=1}^{3} \frac{N_j(e_{g}^{(0)}, e_{g}^{(1)}, e_{g}^{(2)}, a^{(1)}, a^{(2)})}{R_j(e_{g}^{(b)}, e_{g}^{(1)}, e_{g}^{(2)}, a^{(1)}, a^{(2)})} \right\}
\end{align*}
\] (6.7)

where

\[
N_j(e_{g}^{(0)}, e_{g}^{(1)}, e_{g}^{(2)}, a^{(1)}, a^{(2)}) = (e_{g}^{(2)} - e_{g}^{(0)}) e_{g}^{(1)} v^{(2)}(a^{(2)}) - (e_{g}^{(1)} - e_{g}^{(0)}) \times \]

\[
\left[ \left[ A_j^{(2)}(\alpha_1^{(2)}, \alpha_2^{(2)}) - 1 \right] e_{g}^{(1)} - A_j^{(2)}(\alpha_1^{(2)}, \alpha_2^{(2)}) e_{g}^{2} \right] v^{(1)}(a^{(1)}, a^{(2)}) \right\} \quad (6.8)
\]

and

\[
R_j(e_{g}^{(b)}, e_{g}^{(1)}, e_{g}^{(2)}, a^{(1)}, a^{(2)}) = \left[ \left[ A_j^{(2)}(\alpha_1^{(2)}, \alpha_2^{(2)}) - 1 \right] e_{g}^{(1)} - A_j^{(2)}(\alpha_1^{(2)}, \alpha_2^{(2)}) e_{g}^{2} \right] \times \]

\[
\left[ \left[ A_j^{(1)}(\alpha_1^{(1)}, \alpha_2^{(1)}) - 1 \right] e_{g}^{(b)} - A_j^{(1)}(\alpha_1^{(1)}, \alpha_2^{(1)}) e_{g}^{1} \right] (a_1^{(1)} a_2^{(1)} a_3^{(1)}) + A_j^{(1)}(\alpha_1^{(1)}, \alpha_2^{(1)}) \times \]

\[
\left[ 1 - A_j^{(1)}(\alpha_1^{(1)}, \alpha_2^{(1)}) \right] (e_{g}^{(b)} - e_{g}^{(1)}) (e_{g}^{(1)} - e_{g}^{(2)}) (a_1^{(1)} a_2^{(1)} a_3^{(1)}) \right\} \quad (6.9)
\]

The volume of the outer confocal shell and inner ellipsoid are

\[
v^{(1)}(a^{(1)}, a^{(2)}) = \frac{4}{3} \pi \left[ [a_1^{(1)} a_2^{(1)} a_3^{(1)}] - [a_1^{(2)} a_2^{(2)} a_3^{(2)}] \right] \quad (6.10)
\]

and

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\[ V^{(2)}(a^{(2)}) = \frac{4}{3} \pi \left( a_1^{(2)} a_2^{(2)} a_3^{(2)} \right), \quad (6.11) \]

respectively.

To obtain an estimate for a homogeneous inclusion possessing surface phenomena, it is necessary to consider the limiting case as \( \xi \to 0 \). Using a Taylor series expansion, the following approximations for quantities pertaining to the outer ellipsoidal shell can be obtained when \( \xi \) is very small relative to \( a_j^{(2)} \) (\( j = 1, 2, 3 \)):

\[ a_j^{(1)} = a_j^{(2)} + \xi \left( 2a_j^{(2)} \right)^{-1} \quad (j = 1, 2, 3), \quad (6.12) \]

\[ a_1^{(1)} a_2^{(1)} a_3^{(1)} = a_1^{(2)} a_2^{(2)} a_3^{(2)} + \xi \, C(a^{(2)}), \quad (6.13) \]

\[ V^{(1)}(a^{(1)}, a^{(2)}) = \frac{4}{3} \pi \left[ \xi \, C(a^{(2)}) + \xi^2 D(a^{(2)}) \right], \quad (6.14) \]

and

\[ A_j^{(1)}(\alpha_1^{(1)}, \alpha_2^{(1)}) = A_j^{(2)}(\alpha_1^{(2)}, \alpha_2^{(2)}) + \xi \, F_j(a^{(2)}) \quad (j = 1, 2, 3), \quad (6.15) \]

where

\[ C(a^{(2)}) = \frac{1}{2} \left( \frac{a_2^{(2)} a_3^{(2)}}{a_1^{(2)}} + \frac{a_1^{(2)} a_3^{(2)}}{a_2^{(2)}} + \frac{a_1^{(2)} a_2^{(2)}}{a_3^{(2)}} \right), \quad (6.16) \]
\begin{equation}
D(a^{(2)}) = \frac{1}{4} \left( \frac{a_{1}^{(2)}}{a_{2}^{(2)} a_{3}^{(2)}} + \frac{a_{2}^{(2)}}{a_{1}^{(2)} a_{3}^{(2)}} + \frac{a_{3}^{(2)}}{a_{1}^{(2)} a_{2}^{(2)}} \right) - \frac{1}{2} \left[ \frac{a_{2}^{(2)} a_{3}^{(2)}}{(a_{1}^{(2)})^3} + \frac{a_{3}^{(2)} a_{1}^{(2)}}{(a_{2}^{(2)})^3} + \frac{a_{1}^{(2)} a_{2}^{(2)}}{(a_{3}^{(2)})^3} \right] \tag{6.17}
\end{equation}

and

\begin{equation}
F_{1}(a^{(2)}) = \frac{1}{2} \left\{ \left[ \sum_{j=1}^{3} (a_{j}^{(2)})^{-2} \right] A_{j}^{(2)}(\alpha_{1}^{(1)}, \alpha_{2}^{(1)}) - \frac{1}{2} (a_{1}^{(2)} a_{2}^{(2)} a_{3}^{(2)}) \times \right. \\
\int_{0}^{\infty} \left[ 3 \left[ s + (a_{2}^{(2)})^2 \right] \left[ s + (a_{3}^{(2)})^2 \right] + \left[ s + (a_{1}^{(2)})^2 \right] \left[ s + (a_{3}^{(2)})^2 \right] + \left[ s + (a_{1}^{(2)})^2 \right] \left[ s + (a_{2}^{(2)})^2 \right] \times \right. \\
\left[ s + (a_{1}^{(2)})^{5/2} \left[ s + (a_{2}^{(2)})^{3/2} \right] \left[ s + (a_{3}^{(2)})^{3/2} \right] \right] \left. \right\} \cdot \tag{6.18}
\end{equation}

Expressions for \( F_{2}(a^{(2)}) \) and \( F_{3}(a^{(2)}) \) are similar to Equation (6.18) with the appropriate permutation of indices.

During the limiting process, it is assumed the total number of the unit dipoles and charge carriers within the outer shell are conserved. Since generalized permittivity is directly proportional to the density of unit dipoles and unit charge carriers, \( \varepsilon_{g}^{(1)} \) is related to an equivalent generalized surface permittivity \( \lambda_{g}^{(1)} \) in the limiting case as follows:

\begin{equation}
\varepsilon_{g}^{(1)} V^{(1)}(a^{(1)}, a^{(2)}) = \lambda_{g}^{(1)} S(a^{(2)}) \tag{6.19}
\end{equation}
where $S(a^{(2)})$ is the surface area of the inner ellipsoid and $\lambda_{g}^{(1)}$ describes the electric nature of the surface phenomena for the homogeneous inclusion. Using Equation (6.14), it is found that for small $\xi$

\[
\varepsilon_{g}^{(1)} = \frac{3\lambda_{g}^{(1)} S(a^{(2)})}{4\pi C(a^{(2)})} \left[ \xi^{-1} + \frac{D(a^{(2)})}{C(a^{(2)})} \right]. \tag{6.20}
\]

By inserting Equations (6.12) - (6.15) and (6.20) into Equation (6.7) and taking the limit $\xi \to 0$, the following approximation is obtained for the expected value of the volume integral over a homogeneous inclusion with principal semi-axis lengths ($a_1, a_2, a_3$) and generalized permittivity $\varepsilon_{g}^{(i)}$ possessing surface phenomena described by the surface permittivity $\lambda_{g}^{(i)}$:

\[
\left( \mathcal{J} \int_{V_{inc}} \left( \varepsilon(x) - \varepsilon_{g}^{(0)} \right) E^{(i)}(x) \mathrm{d}v \right)_{V_{inc}} = \frac{1}{3} E_{a} \varepsilon_{g}^{(b)} V_{inc} \left\{ \sum_{j=1}^{3} \frac{M_{j}(\varepsilon_{g}^{(0)}, \varepsilon_{g}^{(i)}, \lambda_{g}^{(i)}, a)}{P_{j}(\varepsilon_{g}^{(b)}, \varepsilon_{g}^{(i)}, \lambda_{g}^{(i)}, a)} \right\} \tag{6.21}
\]

where

\[
M_{j}(\varepsilon_{g}^{(0)}, \varepsilon_{g}^{(i)}, \lambda_{g}^{(i)}, a) = \varepsilon_{g}^{(i)} - \varepsilon_{g}^{(0)} + [1 - A_{j}(\alpha_{1}, \alpha_{2})] S(a) V_{inc} \lambda_{g}^{(i)} \tag{6.22}
\]

and

\[
P_{j}(\varepsilon_{g}^{(b)}, \varepsilon_{g}^{(i)}, \lambda_{g}^{(i)}, a) = A_{j}(\alpha_{1}, \alpha_{2}) \varepsilon_{g}^{(i)} + [1 - A_{j}(\alpha_{1}, \alpha_{2})] \varepsilon_{g}^{(b)} + 3A_{j}(\alpha_{1}, \alpha_{2}) \times
\]

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\[ S(a) \left[ F_j(a) + [1 - A_j(\alpha_1, \alpha_2)] C(a) (a_1 a_2 a_3)^{-1} \right] [4\pi C(a)]^{-1} \lambda_g^{(i)}. \] (6.23)

Unlike the result obtained in Chapter Three for an inclusion without surface phenomena, it is apparent that the dielectric response of this inclusion is explicitly dependent on its dimensions. In addition, it can be seen that this approximation differs significantly from that proposed by O'Konski (1960); the effects of the surface phenomena are not expressed simply in terms of an effective inclusion permittivity.

In order to test the validity of this approach, let us briefly consider two specific cases for the approximation given by Equation (6.21). First, for an ellipsoidal inclusion without surface phenomena (i.e. \( \lambda_g^{(i)} = 0 \)), the result derived in Chapter Three for a homogeneous inclusion is obtained. Second, in the case of a spherical inclusion, Equation (6.21) gives

\[
\int_{V_{inc}} \left( \varepsilon_g(x) - \varepsilon_g^{(0)} \right) E^{(i)}(x) \, dv = 3 \varepsilon_a \, V_{inc} \, \frac{\left( \varepsilon_g^{(i)} + 2r^{-1} \lambda_g^{(i)} \right) - \varepsilon_g^{(0)}}{\left( \varepsilon_g^{(i)} + 2r^{-1} \lambda_g^{(i)} \right) + 2\varepsilon_g^{(0)}}. \] (6.24)

This result is identical to the form obtained by O'Konski (1960) and others using boundary conditions to explicitly incorporate surface phenomena.

The form of this approximation can be greatly simplified by restricting the inclusion shape to a spheroid of rotation. Let \( a = a_1 = a_2 \) and \( b = a_3 \); the shape of the inclusion is uniquely defined by its aspect ratio, \( \alpha = b/a \). In this case,

\[ F_j(a) = \frac{1}{4} \left[ (2a^2 + b^2) (1 - A(\alpha)) - 2a^2 \right] \quad (j = 1, 2), \] (6.25)

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\( F_j(a) = \frac{1}{2} \left[ (2a^2 + b^2) A(\alpha) - b^2 \right] \quad (j = 3), \quad (6.26) \)

and

\[ C(a) = \frac{1}{2} \left( 2b + \frac{a^2}{b} \right), \quad (6.27) \]

where \( A(\alpha) \) is the depolarization coefficient associated with the unique axis. Incorporating Equations (6.25)-(6.27) into Equation (6.21), the following estimate is obtained for the expected value of the volume integral in the case of a spheroidal inclusion:

\[
\left\langle \int_{V_{\text{inc}}} \left( \epsilon_g(x) - \epsilon_g^{(0)} \right) E^{(i)}(x) \, dv \right\rangle = \frac{1}{3} E_a \epsilon_g^{(b)} V_{\text{inc}} \times
\]

\[
\left\{ \frac{4 \epsilon_g^{(i)} - \epsilon_g^{(0)}}{[1 - A(\alpha)] \epsilon_g^{(i)} + [1 + A(\alpha)] \epsilon_g^{(b)} + [1 - A(\alpha)] (\alpha^2 + 1) (2\alpha^2 + 1)^{1/2} \beta(b, \alpha) \lambda_g^{(i)}}{[1 - A(\alpha)] \epsilon_g^{(i)} + [1 + A(\alpha)] \epsilon_g^{(b)} + 2A(\alpha) \alpha^2 (2\alpha^2 + 1)^{1/2} \beta(b, \alpha) \lambda_g^{(i)}} \right\}. \quad (6.28)
\]

The size dependence of the inclusion response is determined by \( \beta(b, \alpha) \), the surface area to volume ratio for the inclusion. For oblate and prolate spheroids, \( \beta \) is given by
\[ \beta(b, \alpha) = \frac{3}{2b} (1 + \alpha(\alpha^2 - 1)^{1/2} \log[(\alpha^2 + 1)^{1/2} + \alpha^{-1}]) \]  

(6.29)

and

\[ \beta(b, \alpha) = \frac{3}{2b} (1 + \alpha(1 - \alpha^2)^{-1/2} \sin^{-1}[(1 - \alpha^2)^{1/2}]) \]  

(6.30)

respectively. Further, the quantity \( \beta(b, \alpha) \lambda_g^{(i)} \) completely characterizes the effect of the surface phenomena for all spheroidal particles having the same shape.

**Numerical Simulation**

To illustrate the effect of surface phenomena on the effective dielectric response of a heterogeneous media, a system containing identical spheroidal particles will be used. Utilizing Equation (6.28) to estimate the perturbation term in Equation (6.2) and employing the Lorentz approximation, the dielectric response of the resulting medium is described by

\[
\frac{\varepsilon_g - \varepsilon_g^{(0)}}{\varepsilon_g + 2\varepsilon_g^{(0)}} = \frac{1}{g} c_{inc} \times 
\]

\[
\left\{ \frac{4(\varepsilon_g^{(i)} - \varepsilon_g^{(0)}) + 2[1 + A(\alpha)] \beta(b, \alpha) \lambda_g^{(i)}}{[1 - A(\alpha)] \varepsilon_g^{(i)} + [1 + A(\alpha)] \varepsilon_g^{(0)} + [1 - A(\alpha)](\alpha^2 + 1)(2\alpha^2 + 1)^{1/2} \beta(b, \alpha) \lambda_g^{(i)}} \right\} 
\]
\[
\begin{align*}
\frac{\varepsilon_g^{(i)} - \varepsilon_g^{(0)} + [1 - A(\alpha)] \beta(b, \alpha) \lambda_g^{(i)}}{A(\alpha) \varepsilon_g^{(i)} + [1 - A(\alpha)] \varepsilon_g^{(0)} + 2A(\alpha) \alpha^2 (2\alpha^2 + 1)^{-1} \beta(b, \alpha) \lambda_g^{(i)}} \quad \text{(6.31)}
\end{align*}
\]

where \( c_{\text{inc}} \) is the volume fraction concentration of the inclusions in the system. Further, it will be assumed that the constituents and surface phenomena possess simple linear electrical rheologies of the form given in Equation (4.16) and

\[
\lambda_g^{(i)} = \lambda_a^{(i)} + \frac{\mu_a^{(i)}}{i\omega}, \quad \text{(6.30)}
\]

where \( \lambda_a^{(i)} \) and \( \mu_a^{(i)} \) are the apparent surface permittivity and conductivity, respectively. The parameters for the system are chosen to represent a suspension of colloidal particles in a weak electrolyte (see Table 6.1). The surface phenomenon is taken to be a simple surface conduction (i.e. \( \lambda_a^{(i)} = 0 \). Farads); hence, the effects of this surface mechanism on \( \varepsilon_g^{*} \) is determined by the quantity \( \beta(b, \alpha)\mu_a^{(i)} \).

Figures (6.1)-(6.4) and Figures (6.5)-(6.8) give the computed value of \( \kappa_a^{*} \) and \( \sigma_a^{*} \) for this system as a function of frequency for particle shapes \( \alpha = 1.000, 0.100, 0.010, \) and 0.001, respectively. In all cases, the fractional volume concentration of the particles was taken to be 0.0005. At this concentration level, it can be shown that identical spheroidal particles in this range of aspect ratios used can have uniformly random orientation without particle overlap. For each value of \( \alpha \), the response of the system is calculated for an increasing series of \( \beta(b, \alpha)\mu_a^{(i)} \) values; for a constant value of \( \mu_a^{(i)} \), this represents a
sequence of models where the particle size scale as expressed in terms of $\beta(b, \alpha)$ is increasing.

The presence of the surface conduction causes a significant enhancement of $\kappa_a^*$ at lower frequencies for the more spherical particles (i.e. $\alpha=1.000$ and 0.100). The magnitude of the effect approaches a limiting value as $\beta(b, \alpha)\mu_a^{(i)}$ increases. For the more oblate particles (i.e. $\alpha=0.010$ and 0.001), the existence of the surface conduction initially causes a reduction of $\kappa_a^*$ in the low frequency range that increases with increasing $\beta(b, \alpha)\mu_a^{(i)}$, however, this trend eventually reverses itself. In the case of $\alpha=0.010$, enhancement of $\kappa_a^*$ similar to that observed for the more spherical particles occurs once $\beta(b, \alpha)\mu_a^{(i)}$ is sufficiently large; conversely, enhancement of $\kappa_a^*$ relative to the no surface conduction case is not observed for $\alpha=0.001$ for the values of $\beta(b, \alpha)\mu_a^{(i)}$ used. The effect of surface conduction on $\sigma_a^*$ is similar in all cases; increasing $\beta(b, \alpha)\mu_a^{(i)}$ results in larger values of $\sigma_a^*$ over the high frequency range. Significant enhancement of $\sigma_a^*$ relative to the conductivity of the weak electrolyte over the lower frequencies occurs only in the case of $\alpha=0.001$.

Conclusions

A model has been presented for incorporating the effects of surface phenomena in the dielectric response of a heterogeneous medium which retains the inherent geometrical structure of this system. This is achieved by initially viewing the inclusion as a two-component confocally-layered ellipsoid where the surface phenomena occur in the outer ellipsoidal shell. To obtain the approximation to a homogeneous inclusion system, the outer shell is taken to be infinitesimally thin while the total numbers of unit charge carriers and unit dipoles are conserved. The resulting estimate of the response of such an inclusion
shows that it is explicitly dependent on the dimensions of the inhomogeneity. This approximation is consistent with the response of an inclusion without surface phenomena and the result obtained for a spherical inclusion where surface phenomena are rigorously introduced through use of the appropriate boundary conditions.

When the shape of the inclusions is restricted to spheroids of rotation, it is found that the dielectric response of the inclusion is dependent on its surface area to volume ratio. In addition, it is shown that the electric nature of the surface phenomena for the inclusions of a given shape is completely defined by the product of this ratio and the generalized surface permittivity. Using the approximation for spheroidal inclusions, the dielectric response of a medium containing identical inclusions possessing a surface conduction has been simulated. It is found that the effect of the surface conduction changes significantly as the shape of the inclusion varies.
<table>
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<th>Inclusion</th>
</tr>
</thead>
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<td>$\kappa_a$</td>
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</tr>
<tr>
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<td>$1. \times 10^{-3}$</td>
<td>0.</td>
</tr>
</tbody>
</table>

**Table 6.1:** Dielectric constants and conductivities used for the constituents of the model colloidal suspension.
Figure 6.1: Computed $\kappa_\ast^\alpha$ as a function of frequency (plotted in log scale) for models containing identical inclusion possessing a surface conduction. Value of $\beta(b, \alpha)\mu_{\ast\alpha}^{(i)}$ used in a given model is denoted. Aspect ratio ($\alpha$) of inclusion used in all models is 1.000.
Figure 6.2: Computed $\kappa^*_a$ as a function of frequency (plotted in log scale) for models containing identical inclusion possessing a surface conduction. Value of $\beta(b, \alpha)\mu_\alpha^{(i)}$ used in a given model is denoted. Aspect ratio ($\alpha$) of inclusion used in all models is 0.100.
Figure 6.3: Computed $\kappa^*_a$ as a function of frequency (plotted in log scale) for models containing identical inclusion possessing a surface conduction. Value of $\beta(b, \alpha)\mu_{(i)}^{(i)}$ used in a given model is denoted. Aspect ratio ($\alpha$) of inclusion used in all models is 0.010.
Figure 6.4: Computed $\kappa^*$ as a function of frequency (plotted in log scale) for models containing identical inclusion possessing a surface conduction. Value of $\beta(b, \alpha)\mu_a^{(i)}$ used in a given model is denoted. Aspect ratio ($\alpha$) of inclusion used in all models is 0.001.
Figure 6.5: Computed $\sigma_a^*$ as a function of frequency (plotted in log scale) for models containing identical inclusion possessing a surface conduction. Value of $\beta(b, \alpha)\mu_0^{(i)}$ used in a given model is denoted. Aspect ratio ($\alpha$) of inclusion used in all models is 1.000.
Figure 6.6: Computed $\sigma_n^*$ as a function of frequency (plotted in log scale) for models containing identical inclusion possessing a surface conduction. Value of $\beta(b, \alpha)\mu_n^{(i)}$ used in a given model is denoted. Aspect ratio ($\alpha$) of inclusion used in all models is 0.100.
Figure 6.7: Computed $\sigma^*_a$ as a function of frequency (plotted in log scale) for models containing identical inclusion possessing a surface conduction. Value of $\beta(b, \alpha)\mu^{(i)}_a$ used in a given model is denoted. Aspect ratio ($\alpha$) of inclusion used in all models is 0.010.
Figure 6.8: Computed $\sigma_a^*$ as a function of frequency (plotted in log scale) for models containing identical inclusion possessing a surface conduction. Value of $\beta(b, \alpha)\mu_a^{(i)}$ used in a given model is denoted. Aspect ratio ($\alpha$) of inclusion used in all models is 0.001.
Chapter Seven

Summary

Microscopic phenomena can significantly affect the macroscopic physical properties of a porous rock. A mathematical model is necessary to establish and analyze the connection between a given microscopic phenomenon and the observed macroscopic behavior. Such theoretical formulations have been derived in this study for the relationship between the elastic and dielectric responses of porous rocks and certain microscopic phenomena.

Contact generation has been suggested as a possible reason for the discrepancy between previous models and experimental data for the functional dependence of elastic wave velocities on confining pressure in unconsolidated materials. A model for the elastic behavior of a random sphere packing that permits the creation of new grain contacts during hydrostatic compression was derived for the first time in this thesis. Using this formulation, it was found that the closure of near-contact gaps having a relatively small mean width (i.e. on the order of 1/1000 of a sphere radius) can significantly affect the elastic behavior of the glass bead packings over the range of confining pressures employed in these experiments.

Over a range of conditions that are of interest in the geophysical measurement of elastic and electromagnetic wave propagation in porous rocks, the behavior of the system is significantly affected by the geometrical configuration of the constituents on a
microscopic scale. Sufficient information about the microstructure of porous rocks is obtained by viewing these systems as a rock matrix into which inclusions representing the individual pore spaces are embedded. Inclusion-based formulations can be employed to determine the macroscopic physical properties of a porous rock when this description of its microstructure is adopted. Major classes of these formulations are defined by the manner in which inclusion interactions are simulated in the analysis; for each class of formulations, there is an associated microstructure that specifies the topological relationships between the various components.

Pore-scale fluid distribution in partially saturated porous rocks is a microscopic phenomenon that is characterized by the geometrical configuration of the constituents within the system. Drastically different theoretical formulations were obtained for various styles of pore-scale fluid distribution while maintaining fixed constituent properties and pore structure. It was determined that the pore fluid configuration within the individual pore spaces and the pore geometries in which saturation conditions are varying are critical factors in determining the elastic and electromagnetic wave propagation in partially saturated rocks. Further, the effects of changes in saturation conditions within a particular pore geometry significantly increase as the aspect ratio of the pores decreases (i.e. as the pores become more crack-like).

The mathematical expressions derived in this study were used to analyze the variation in the functional relationships observed between the level of water saturation and the physical properties of a tight gas sandstone and a gas bead packing when both of these systems were subjected to two differing saturation methods. Using simple models that consider only the basic geometrical elements that describe the pore-scale fluid
distribution, good agreement is obtained between the theoretical models and the experimental data. Further, it is demonstrated that a single geometrical model for the pore-scale fluid distribution can be used to simultaneously estimate the dielectric and elastic response of a partially saturated rock. However, this formulation assumes that communication between individual pores is minimal and intrapore phenomena control the physical behavior of the system. This type of formulation does not accurately predict the functional relationships observed in experimental data where these conditions are not satisfied.

Surface phenomena, such as surface conduction and electrical double layers, can have a significant effect on the electrical properties of porous rocks; however, these mechanisms are explicitly excluded in the derivation of the inclusion-based formulations used to describe these systems. Further, incorporation of surface phenomena into the analysis of non-spherical inclusion shapes is very difficult. To obtain an approximation for the effects of this mechanism on the electrical response of an ellipsoidal inclusion, a model based on the limiting case of a confocally-layered inclusion was derived by assuming the conservation of unit charge carriers and unit dipoles. This approach retains the geometrical configuration inherent in the system. The behavior of the inclusion with surface phenomena is found to depend on the absolute size of the inclusion. Using this model, it is shown that the effects of the presence of a surface conduction changes significantly as the shape of the inclusion varies.

The results of this study clearly demonstrate the usefulness of mathematical models in describing the effects of a given microscopic phenomenon on the macroscopic physical
properties of a porous rock. It is significant to note that such an analysis does not always require a detailed description of the porous rock system; good agreement between the observed behavior and a theoretical formulation can be obtained by utilizing a relatively simple model. However, much basic work still remains to be done in this field. Not only do models for other microscopic phenomena and macroscopic physical properties need to be developed, the use of these formulations to rigorously analyze experimental measurements in the context of inverse problem theory must be further explored. Therefore, I foresee much significant work occurring in this field over both the near-term and distant future.
References


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Appendix One

Constitutive Relationships and Generalized Permittivity

The electromagnetic behavior of a homogeneous, isotropic medium is generally expressed in terms of its dielectric permittivity ($\varepsilon$), electrical conductivity ($\sigma$), and magnetic permeability ($\mu$) which describe the capacitive, conductive, and inductive properties, respectively, of the substance. For materials exhibiting a linear electromagnetic behavior, these quantities are defined by the following relationships for the time harmonic fields in complex notation:

$$D = \varepsilon E,$$  \hspace{1cm} (A1.1)

and

$$J = \sigma E,$$  \hspace{1cm} (A1.2)

$$B = \mu H,$$  \hspace{1cm} (A1.3)

where $B$, $D$, $E$, $H$, and $J$ are the magnetic induction, electric displacement, electric field strength, magnetic field strength, and electric free current density, respectively. The quantities $\varepsilon$, $\sigma$, and $\mu$ are the Fourier transforms of the impulse response for each of the
respective properties. In general, \( \varepsilon, \sigma, \) and \( \mu \) are complex functions of the form

\[
\varepsilon = \varepsilon' - i \varepsilon'', \tag{A1.4}
\]

\[
\sigma = \sigma' - i \sigma'', \tag{A1.5}
\]

and

\[
\mu = \mu' - i \mu''. \tag{A1.6}
\]

where the superscripts \( ' \) and \( '' \) denote the real and imaginary parts. Assuming causality, the real and imaginary parts are related through the Hilbert transform.

The electrical properties are generally measured by observing the electric current density which results from an applied electric field. Assuming that the effects of magnetic induction may be neglected, this observed current density is the total current density \( J_t \) where

\[
J_t = J_d + J_f \tag{A1.7}
\]

and

\[
J_d = i \omega D \tag{A1.8}
\]
is the displacement current density associated with the capacitive behavior of the medium (Fuller and Ward, 1970). Therefore, it is useful to revise the constitutive relationship given by Equation (A1.2) to the form

\[ J_t = \sigma_g E, \]  
(A1.9)

where

\[ \sigma_g = \sigma + i\omega \epsilon \]  
(A1.10)

is the generalized electrical conductivity. Conversely, the generalized dielectric permittivity \( \epsilon_g \) can be defined in term of the generalized electric displacement \( D \) as follows:

\[ D = \epsilon_g E \]  
(A1.11)

where

\[ D = D + (i\omega)^{-1} J_t \]  
(A1.12)

and

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\[ \epsilon_g = \epsilon - i \omega^{-1} \sigma. \]  \hspace{1cm} (A1.13)

A complete description of the electrical properties of a linear material can be accomplished by employing either generalized constitutive relationship given by Equations (A1.9) and (A1.11).

The real part of \( \epsilon_g \) and \( \sigma_g \) are referred to as the apparent dielectric permittivity \( \epsilon_a \) and the apparent electrical conductivity \( \sigma_a \), respectively. It is these quantities that are generally reported as the experimentally measured dielectric permittivity and electrical conductivity. Further, coupling of the capacitive and conductive properties occurs at the discrete interfaces within a heterogeneous medium. Hence, generalized constitutive relationship provide an effective means for incorporating both types of phenomena into the inclusion-based formulation for the electrical properties of porous rocks.
Appendix Two

Evaluation of the Depolarization Coefficient Associated with the Unique Principal Axis of the Inner Spheroid within a Confocally-Layered Inclusion

Evaluation of the expression describing the dielectric response of a confocally-layered inclusion requires the determination of \( B(s^{(1)}, \alpha) \), the depolarization coefficient associated with the unique axis of the inner spheroid. For a non-degenerate spheroidal pore shape (i.e. \( 0 < \alpha < 1 \)), this is accomplished by determining the aspect ratio of the inner spheroid (\( \beta \)). Let \( (a_1, a_2) \) and \( (b_1, b_2) \) be the major and minor semi-axis length pair for the pore space and inner spheroid, respectively; then,

\[
\alpha = \frac{a_2}{a_1}
\]  

(A2.1)

and

\[
\beta = \frac{b_2}{b_1}
\]  

(A2.2)

Since the two spheroids are confocal, then the semi-axis lengths are related by

\[
(a_1)^2 - (b_1)^2 = (a_2)^2 - (b_2)^2
\]  

(A2.3)
Using Equations (A2.1) and (A2.2), Equation (A2.3) can be rewritten as

\[ \frac{1-\beta^2}{1-\alpha^2} = \left(\frac{a_1}{b_1}\right)^2. \]  

(A2.4)

The fractional volume of the non-wetting pore fluid phase within the inclusion is given by

\[ s^{(1)} = \frac{b_2 \left(\frac{b_1}{a_1}\right)^2}{a_2 \left(\frac{a_1}{b_1}\right)^2}. \]

(A2.5)

Again using Equations (A2.1) and (A2.2), Equation (A2.5) can be expressed as

\[ \beta^{2/3} = (\alpha s^{(1)})^{2/3} \left(\frac{a_1}{b_1}\right)^2. \]

(A2.6)

Combining Equations (A2.4) and (A2.6), one obtains

\[ \gamma^3 + \left[ \frac{1-\alpha^2}{(\alpha s^{(1)})^{2/3}} \right] \gamma - 1 = 0, \]

(A2.7)

where \( \gamma = \beta^{2/3} \).

Equation (A2.7) is the normal form of a cubic system which is discussed in Burington (1965). For a non-degenerate oblate spheroidal pore, this equation possesses only one real-valued root that is given by

\[ \beta = \left[ \frac{2(\xi_{(l)})^{1/2}}{(3)} \cot(2\phi) \right]^{3/2}, \]

(A2.8)
where

$$\phi = \tan^{-1} \left[ (\tan \varphi)^{1/3} \right], \quad \text{(A2.9)}$$

$$\varphi = \frac{1}{2} \tan^{-1} \left[ \left( \frac{4c^3}{27} \right)^{1/2} \right], \quad \text{(A2.10)}$$

and

$$c = \left[ \frac{1 - \alpha^2}{(\alpha \cdot s^{(1)})^{2/3}} \right]. \quad \text{(A2.11)}$$

Once $\beta$ is determined, $B(s^{(1)}, \alpha)$ is calculated using Equations (4.1) and (4.3). In the degenerate cases of the sphere and the flat disc, $B(s^{(1)}, \alpha)$ is taken to be $\frac{1}{3}$ and 1, respectively.